

# Seismology and Tsunami Warning Center Operations

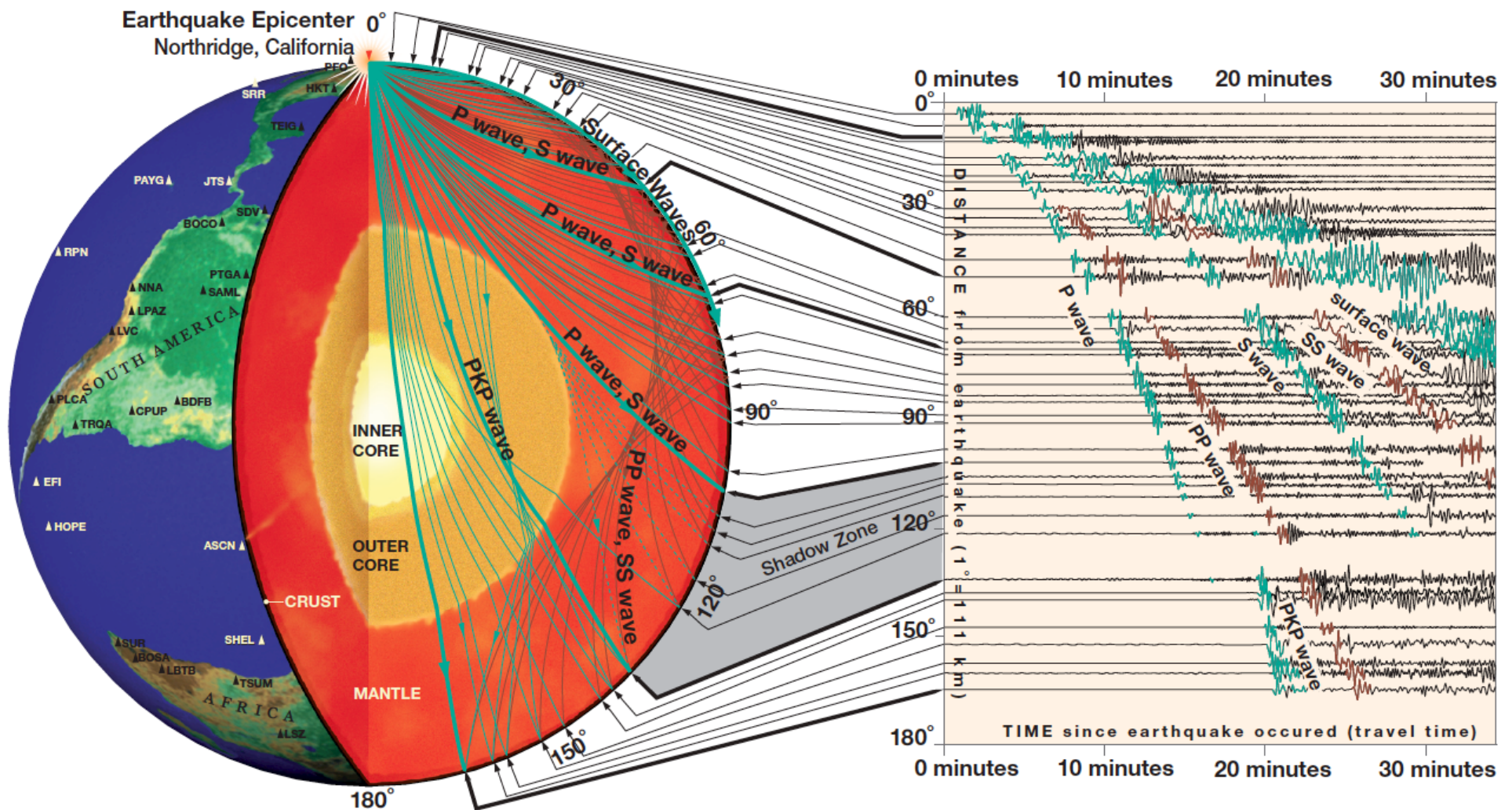
Gerard Fryer  
Pacific Tsunami Warning Center

ITIC Training Program Hawaii  
August 25, 2014

# Earthquake Detection and Characterization

- Motivation: Seismic waves travel roughly 60 times faster than a tsunami, so you can learn a huge amount about an earthquake, including its tsunamigenic potential, before you have any direct information about the tsunami.
- What we solve for now:
  1. Latitude, longitude, depth, origin time (5-7 minutes)
  2. Magnitude (6-8 minutes)
  3. “Fast” or “slow” earthquake? (7-8 minutes)
  4. Mechanism (25 minutes)
- Not yet implemented but high on the wish list:
  1. Magnitude correct to 0.2 within 5 minutes
  2. Rupture direction and extent (finite fault solution)

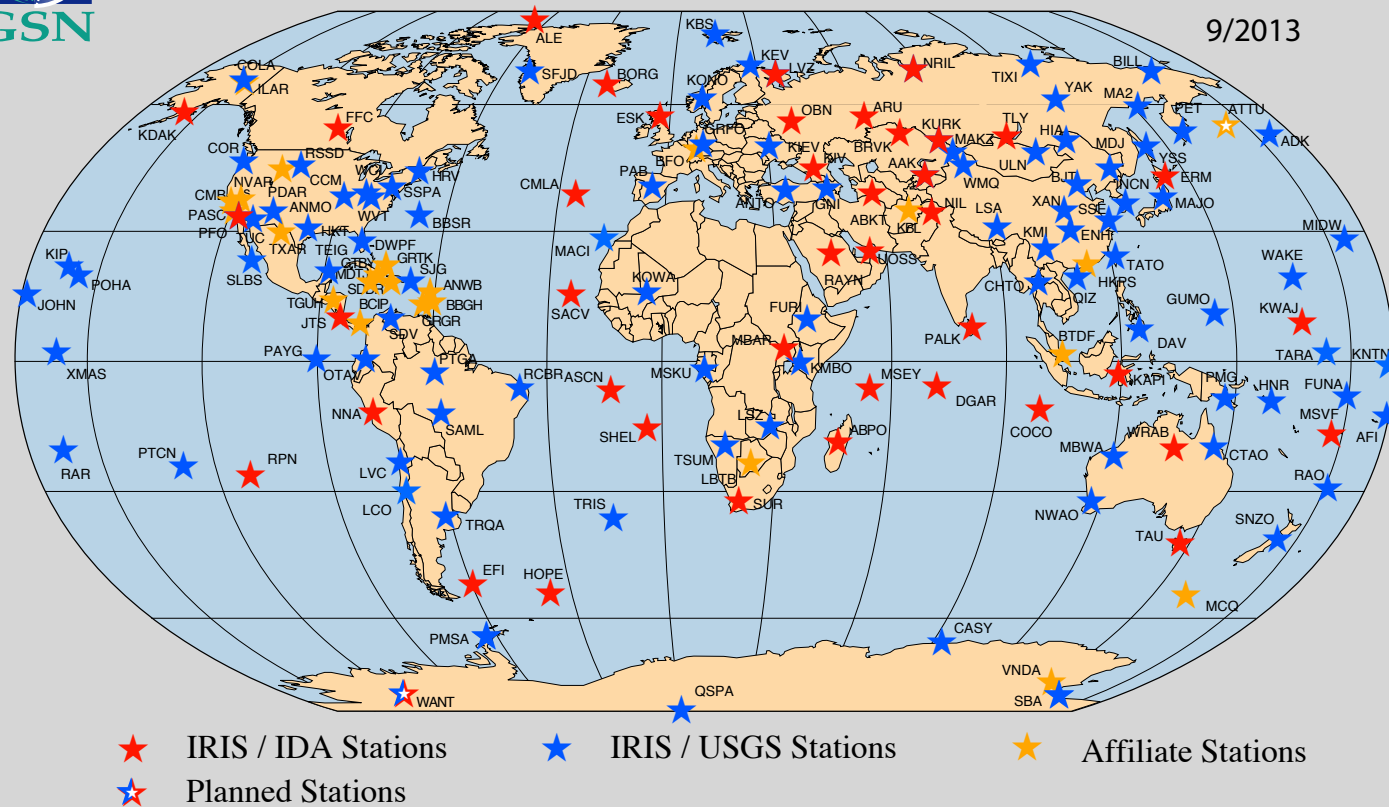
# Seismic travel times



# Seismic Networks



## GLOBAL SEISMOGRAPHIC NETWORK



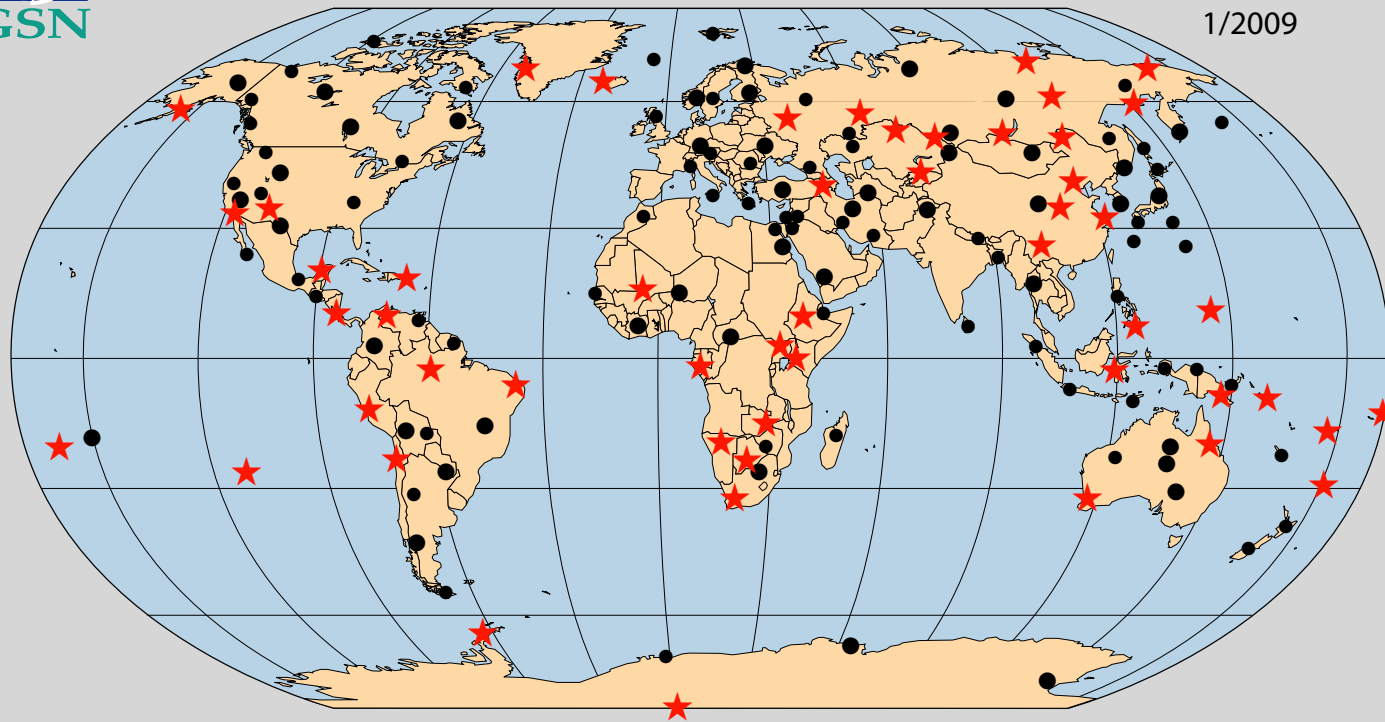


# Seismic Networks



## GLOBAL SEISMOGRAPHIC NETWORK & INTERNATIONAL MONITORING SYSTEM (IMS)

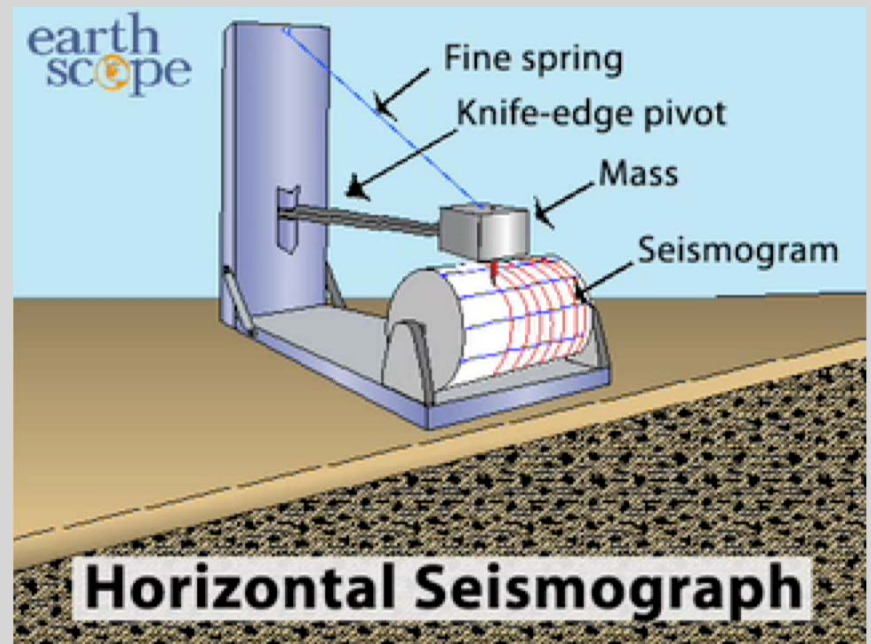
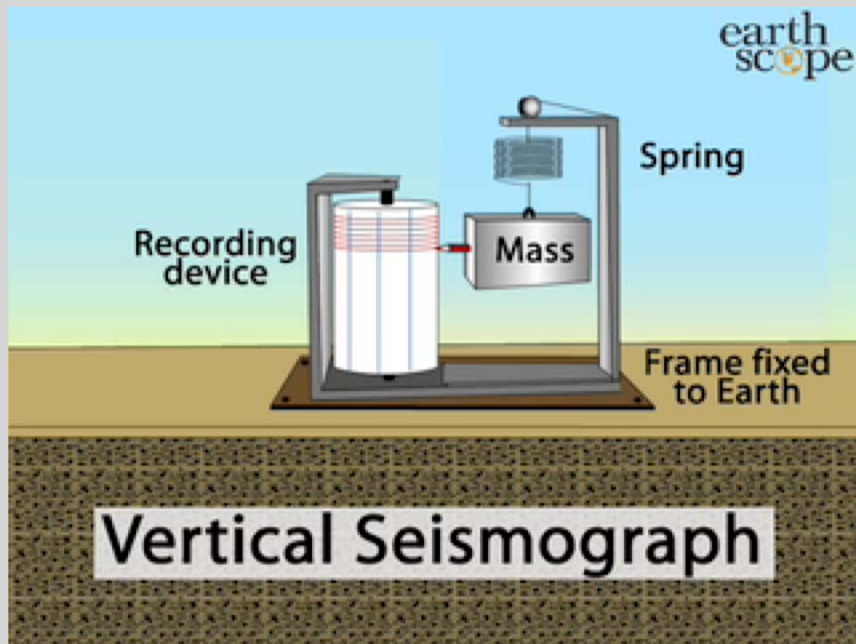
1/2009



- ★ GSN IMS Designated Stations
- Other IMS Seismic Stations

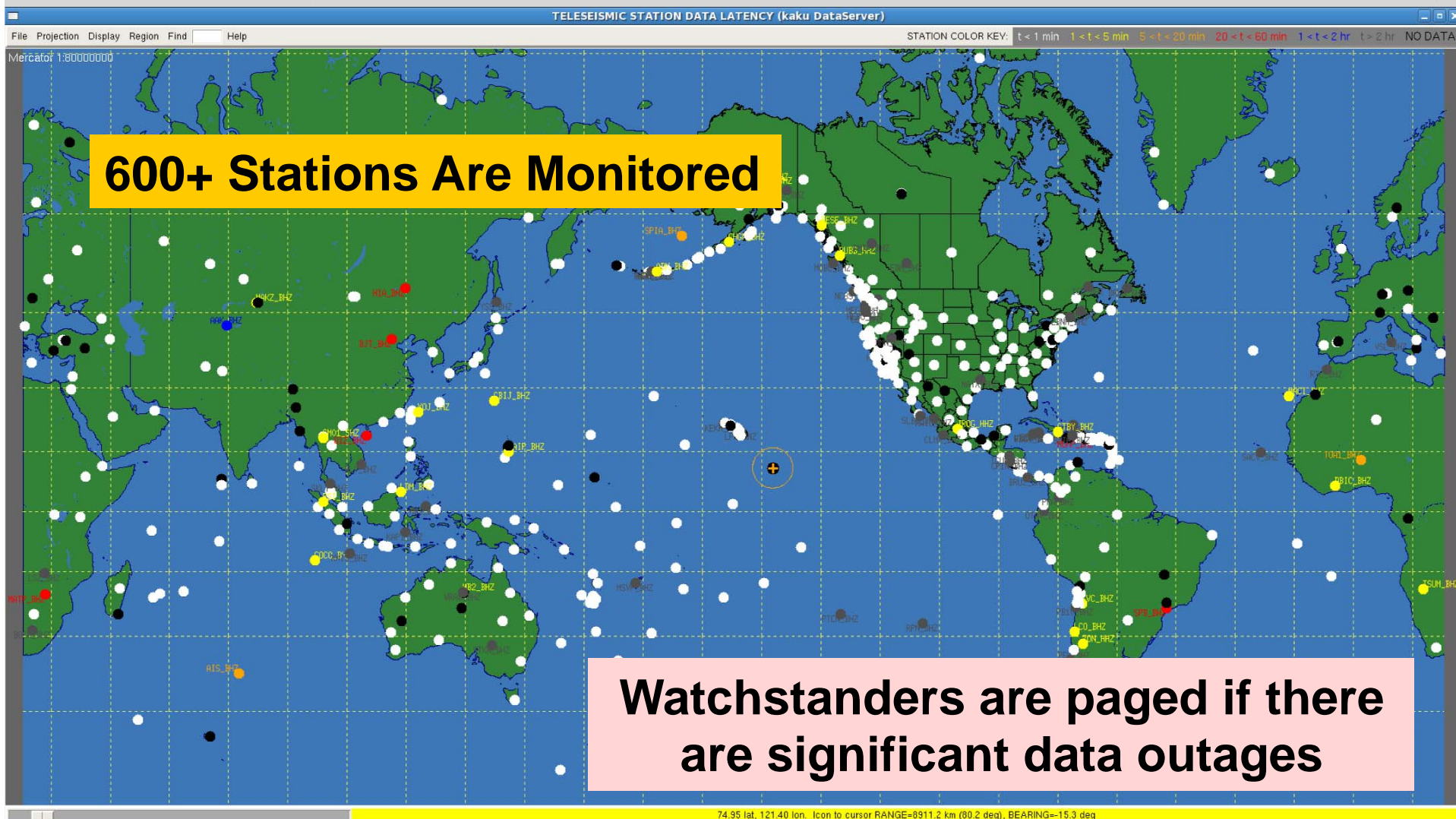
# Seismometers

- We get data from about 650 seismometers worldwide in near-real time (latency is typically a few tens of seconds).
- These are **broadband** seismometers (i.e., capable of recording energy from 1 to 300 seconds period). It is broadband seismometers that make tsunami warning possible.



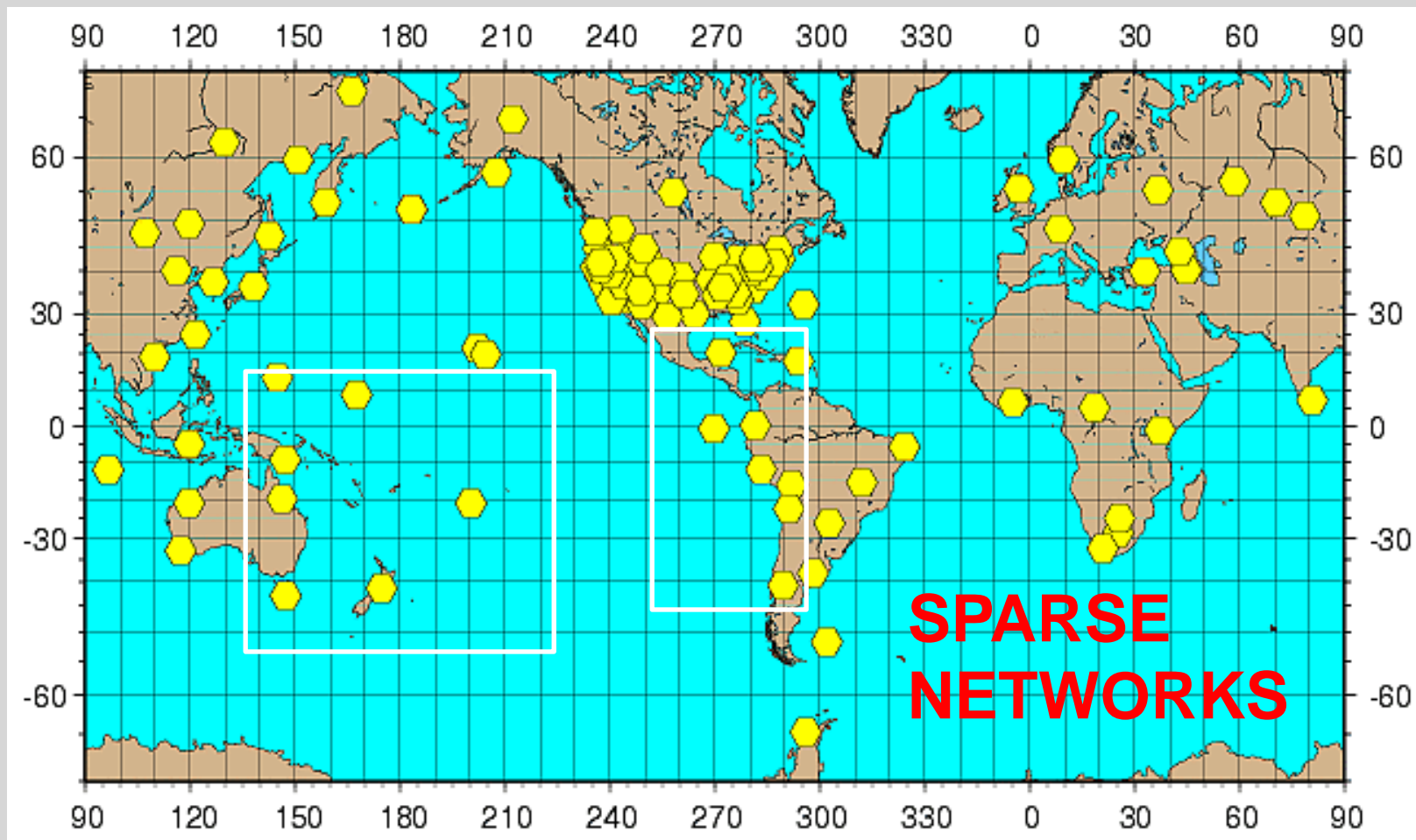
# Data Availability/Latency Monitoring

STATION COLOR KEY:  $t < 1$  min    $1 < t < 5$  min    $5 < t < 20$  min    $20 < t < 60$  min    $1 < t < 2$  hr    $t > 2$  hr   NO DATA



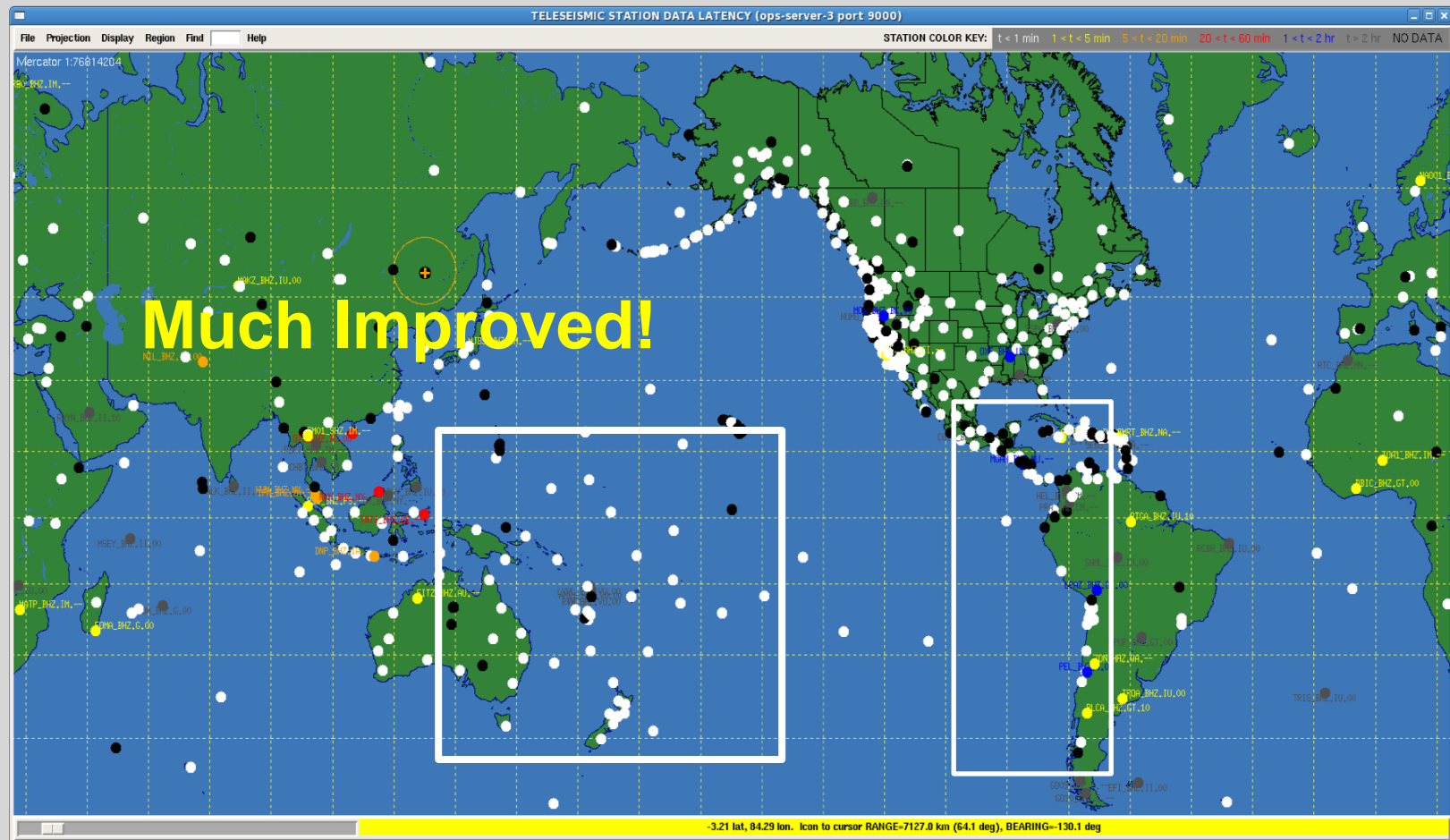


# Broadband Seismometer Distribution Oct. 2007



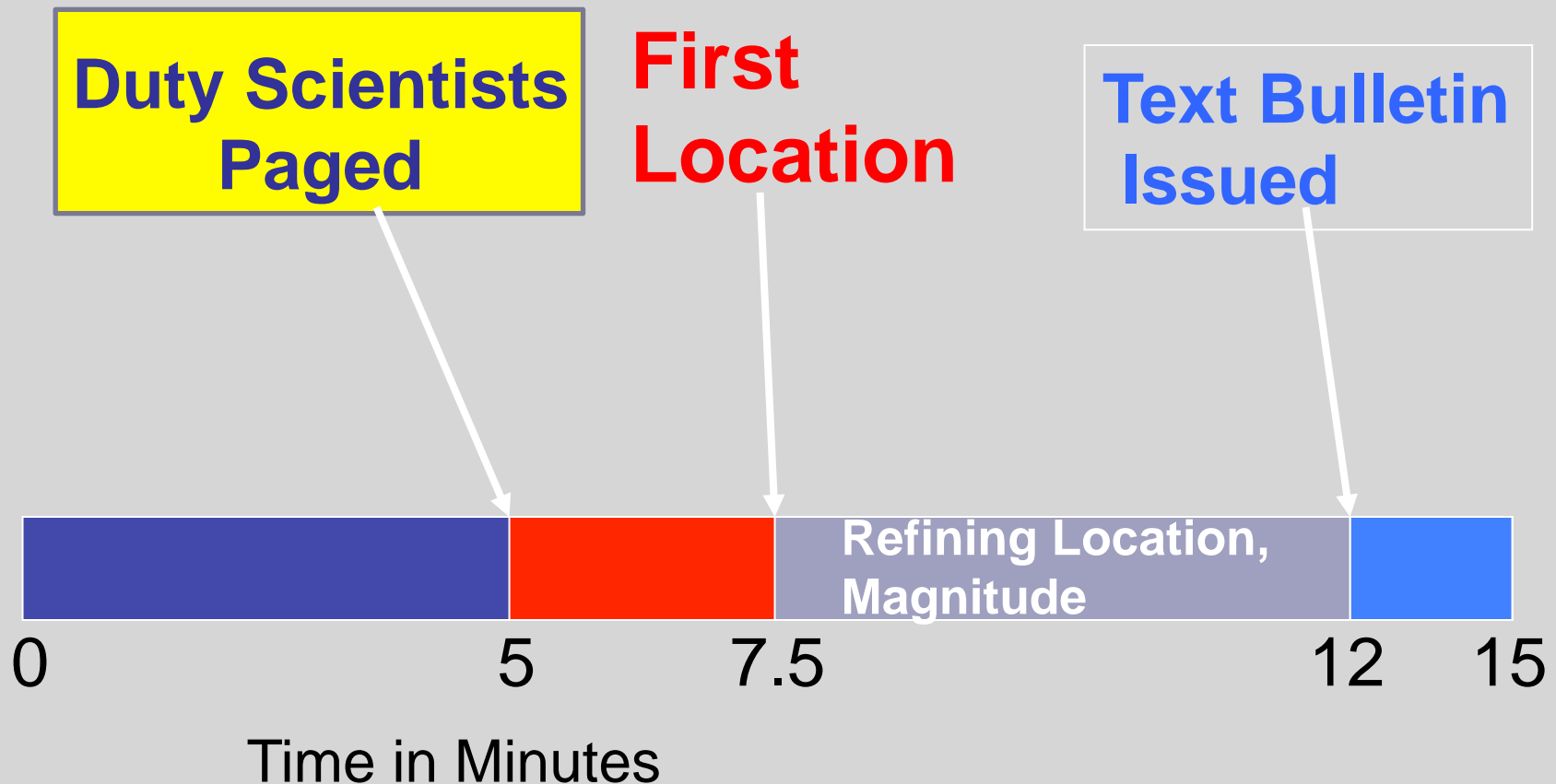


# Broadband Seismometer Distribution May 2014



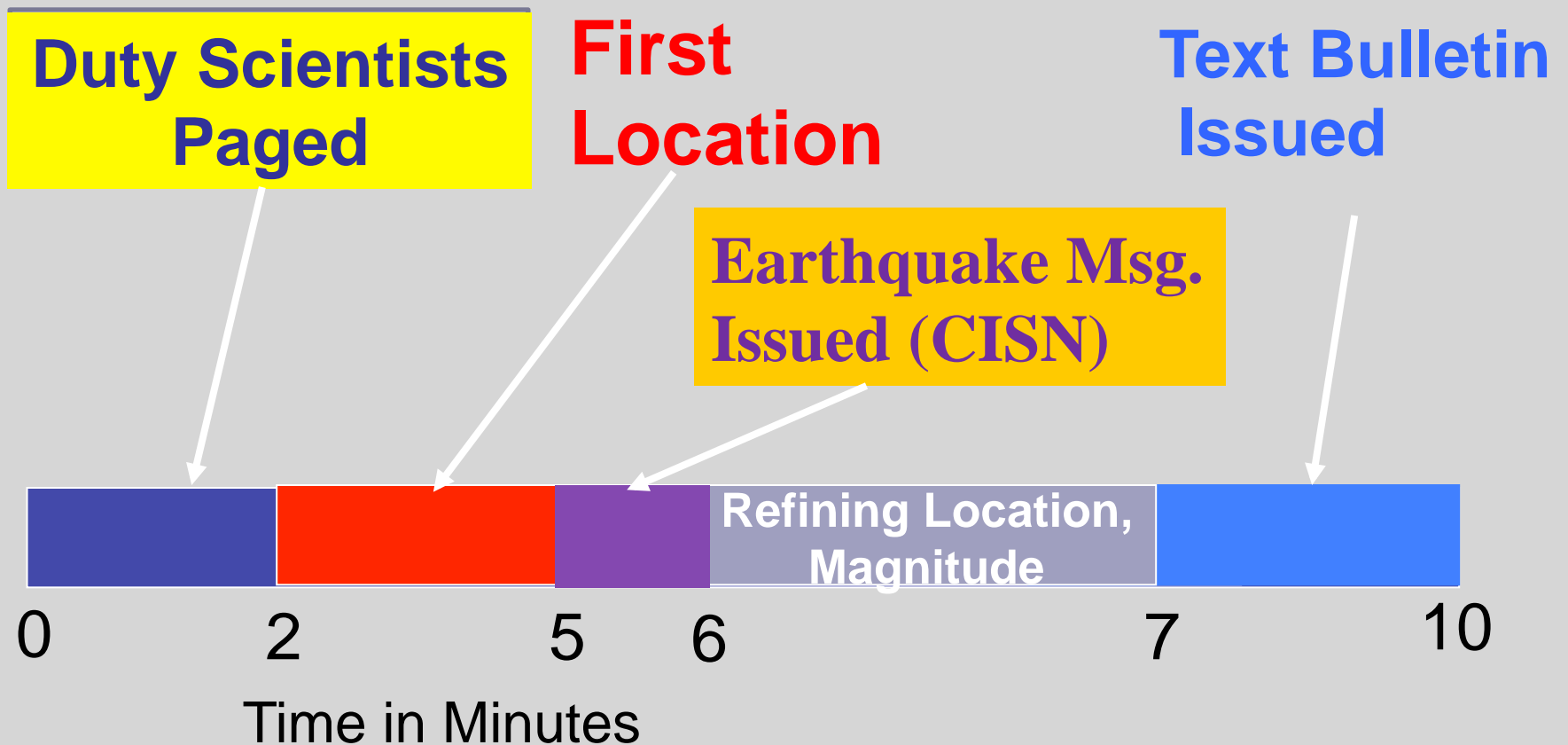
# Earthquake Outside Hawaii Region

## Timeline to Issue Initial Bulletin 2008



# Earthquake Outside Hawaii Region

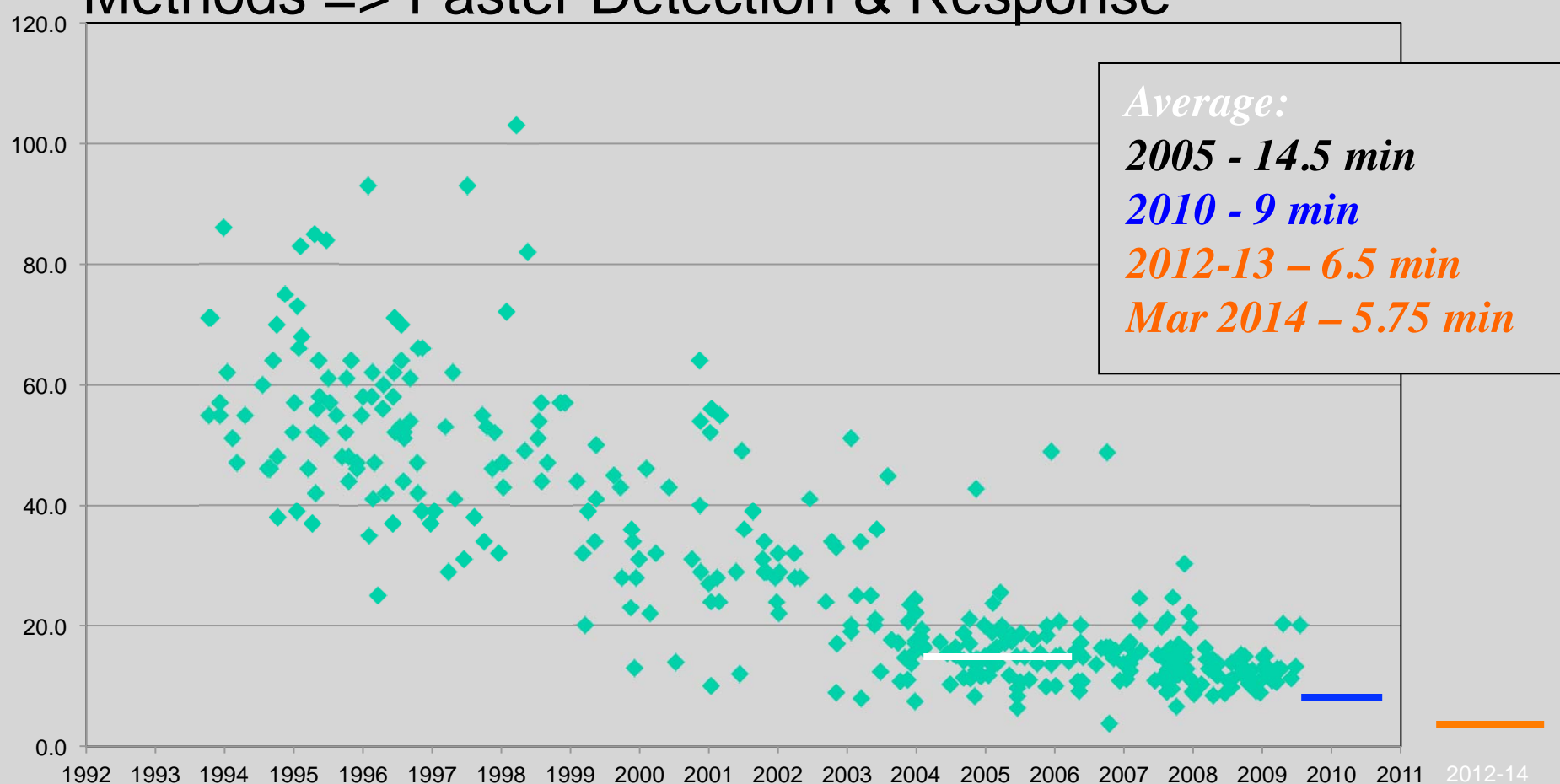
## Timeline to Issue Initial Bulletin 2014





# PTWC Elapsed Time to Initial Bulletin

Earthquakes: Denser  
Networks, Higher Quality Data, Better and Faster  
Methods => Faster Detection & Response



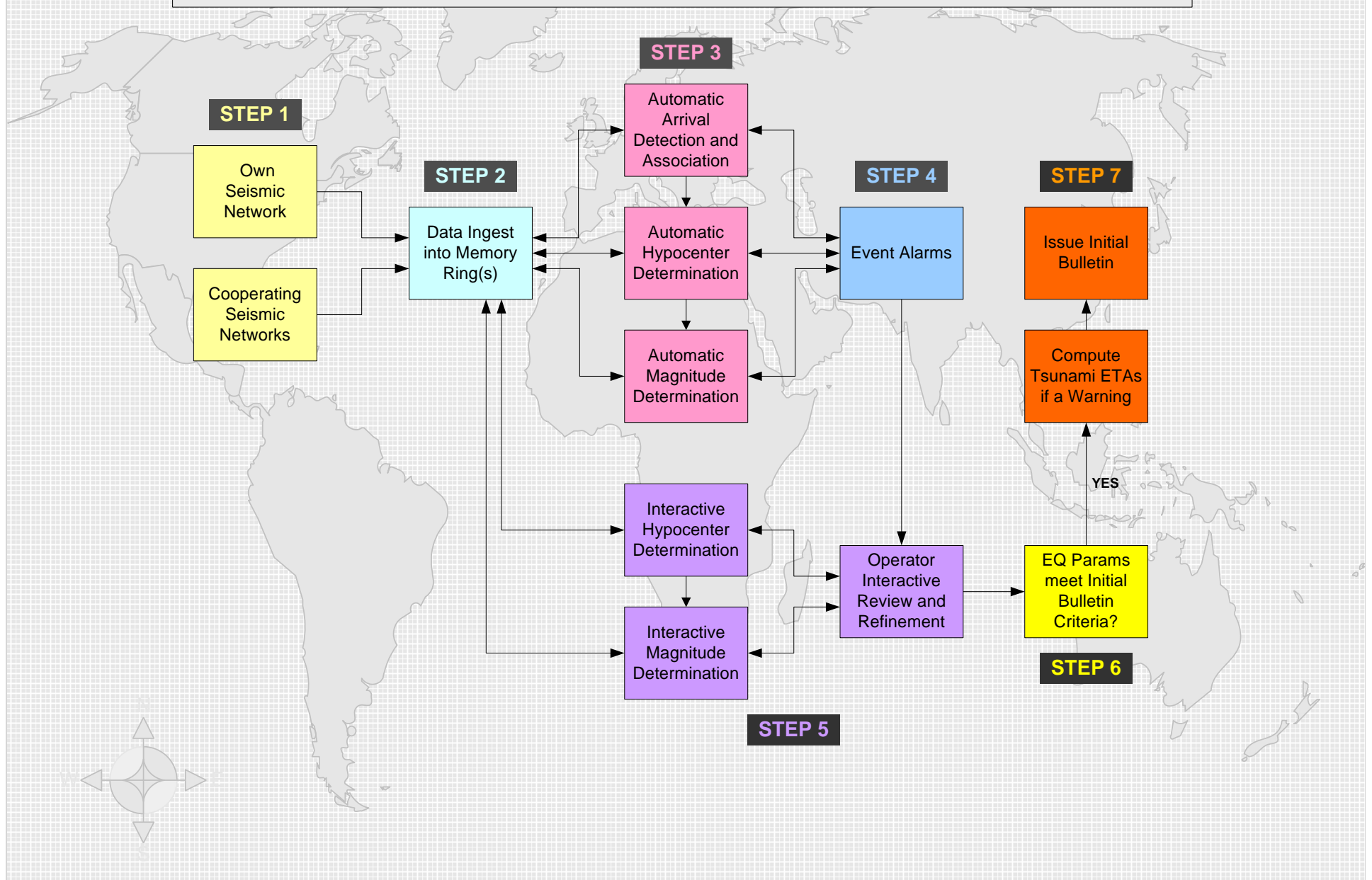


# Current PTWC Products

- Text product only
- Product – Countries put in Warning / Watch status. Estimated arrival times provided

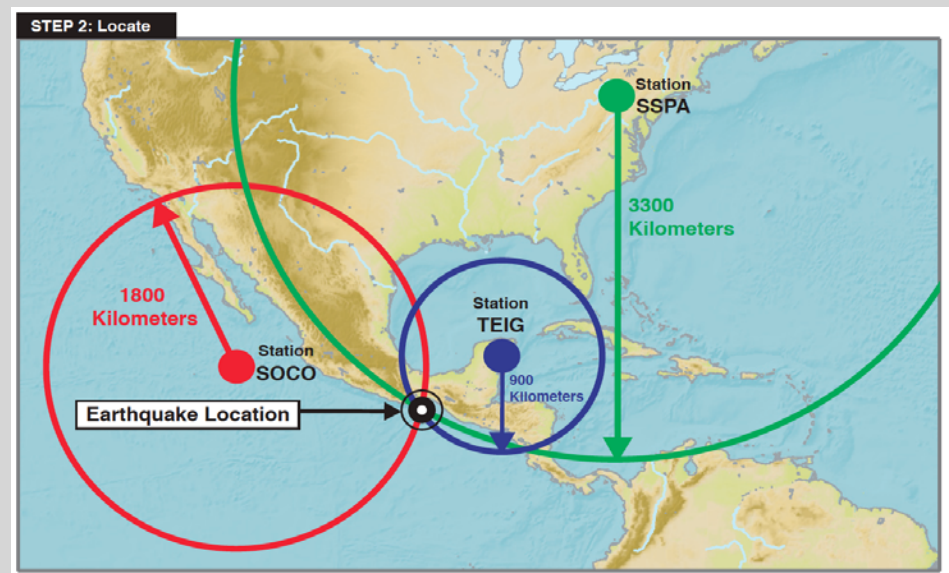
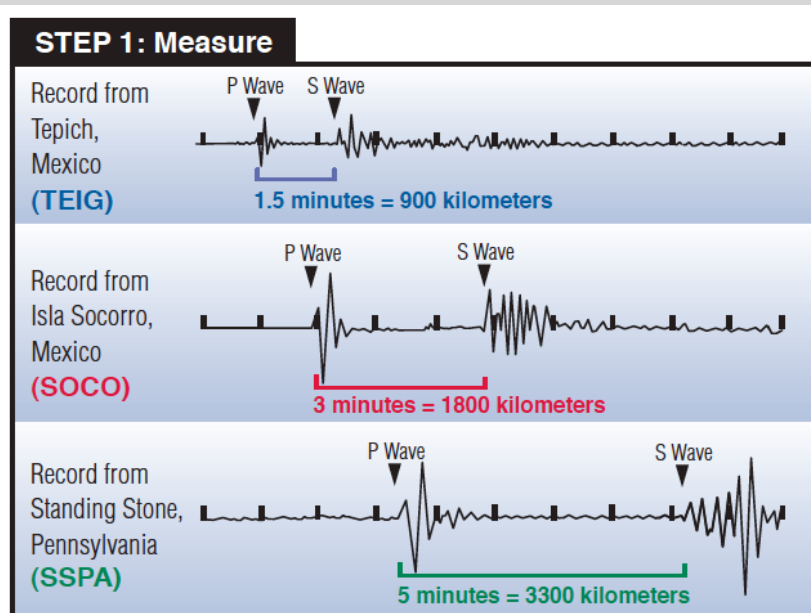
| Earthquake  | Product Type         | Description  |
|---|----------------------|--|
| Magnitude 6.5-7.5 or >6.5 and >100km depth or far inland      | Information Bulletin | No Tsunami Threat or only small Local Tsunami Threat         |
| Magnitude 7.6-7.8 and <100 km depth and under sea or near sea | Regional Warning     | Warning to 1000km from epicenter                             |
| Magnitude >7.8 and <100km depth and under sea or near sea     | Expanding Warning    | Warning if < 3 hours to impact, Watch if 3-6 hours to impact |
| Confirmed major tsunami                                       | Pacific-wide Warning | Confirmed tsunami with widespread destructive threat         |

## PTWC General Processes and Procedures for Initial Tsunami Bulletins



# Locating earthquakes

- Four unknowns: lat, long, depth, origin time.
- Need a minimum of four stations. More is better. The stations may not be well located relative to the earthquake, however, so we often have to fix depth.



# Automation tools

- A *Picker* scans each seismogram and identifies possible seismic arrivals.
- An *Associator* groups picks from different seismograms to try to identify earthquakes.
- If picks are consistent with an earthquake, standard magnitude algorithms compute an automatic magnitude.
- *Paging*:
  - If one seismometer has a pick with a large signal-to-noise ratio, the duty scientist gets a page.
  - If two or more seismometers in a given region have picks within a few minutes of each other, the duty scientist and the standby scientist both get pages.
  - If the automatic magnitude exceeds 5.6, both scientists get pages.



# Automatic Processing

## STEP 2

Data Ingest  
into Memory  
Ring(s)

## STEP 3

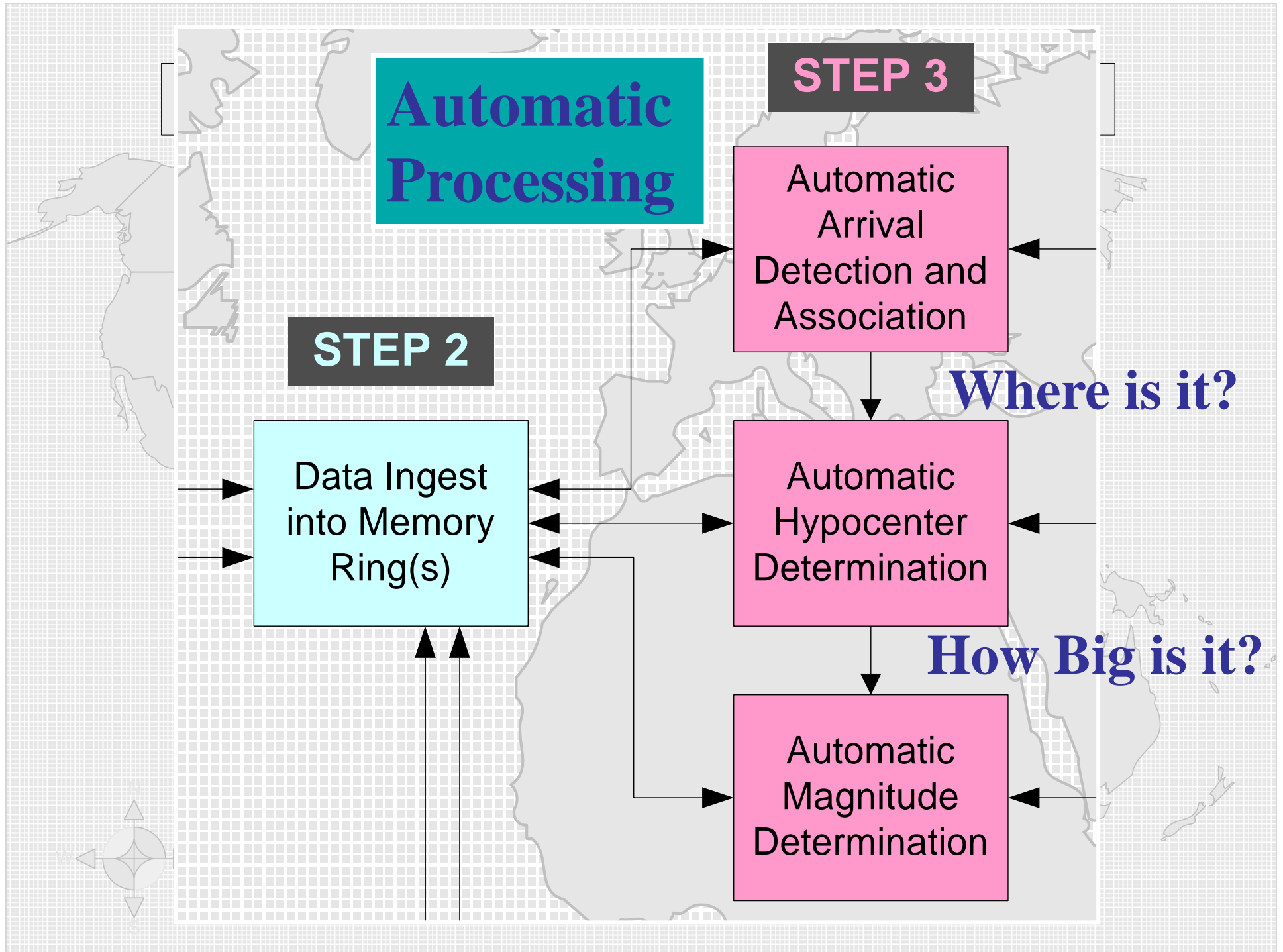
Automatic  
Arrival  
Detection and  
Association

Where is it?

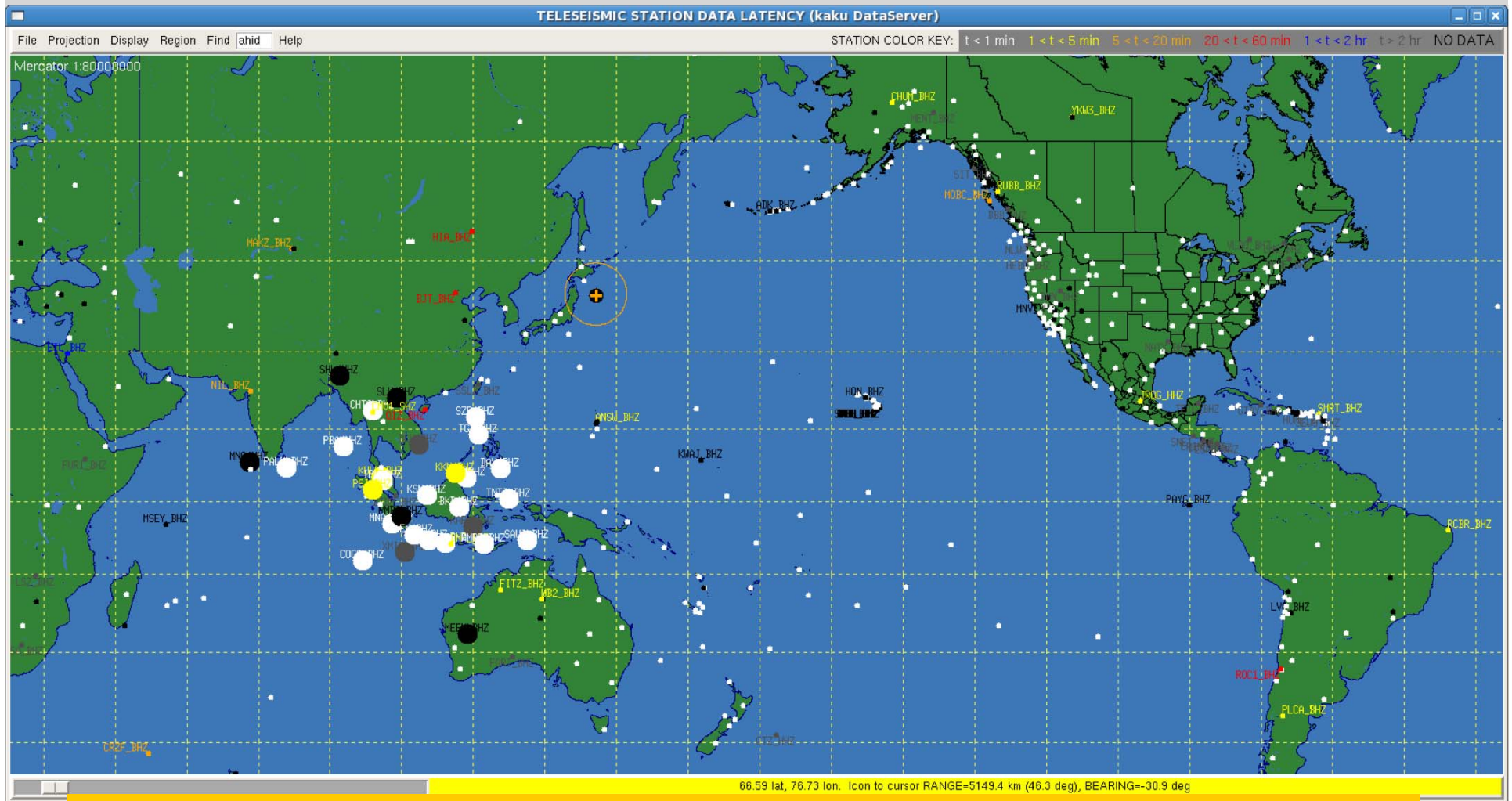
Automatic  
Hypocenter  
Determination

How Big is it?

Automatic  
Magnitude  
Determination



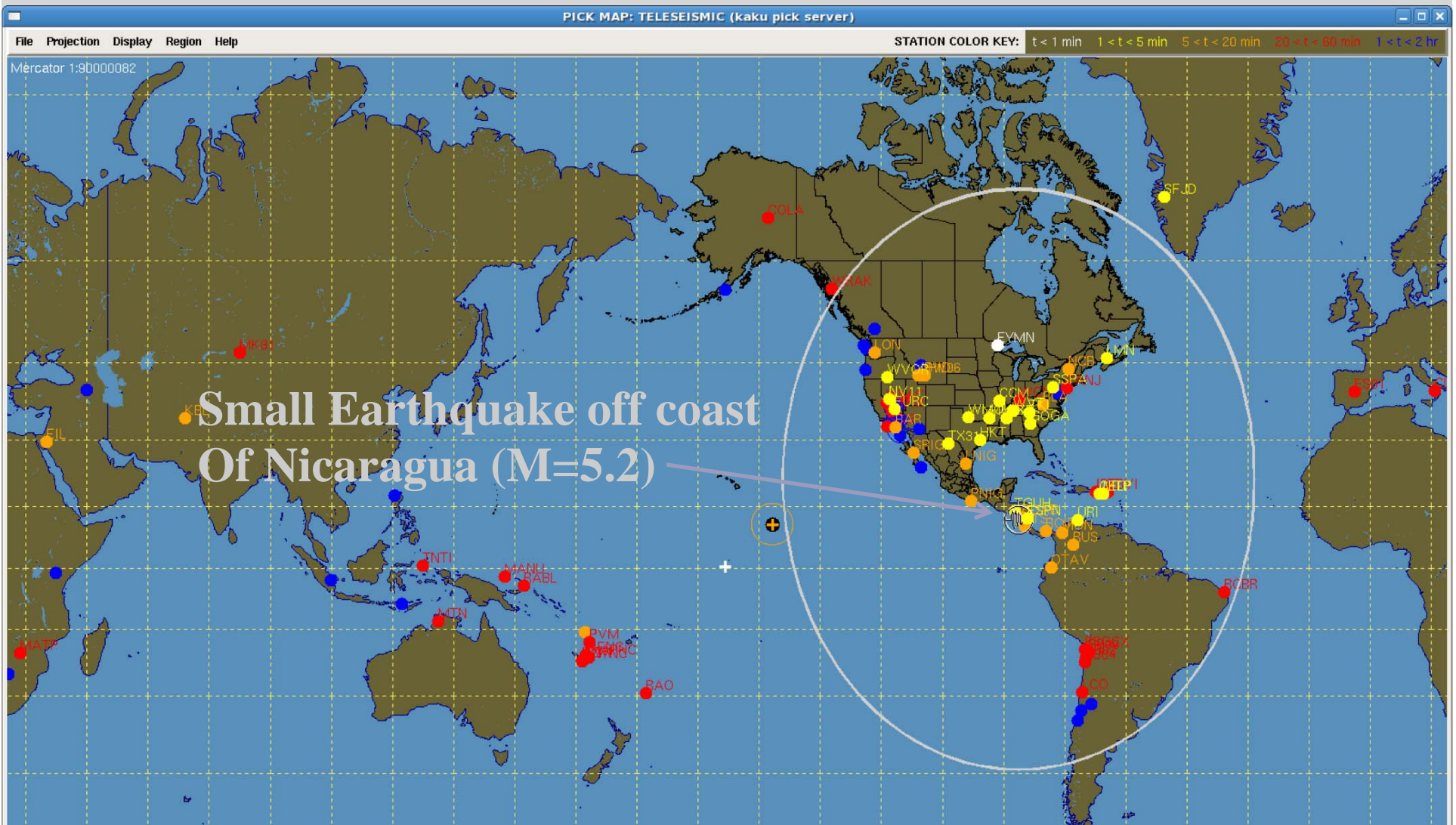
# Broadband Seismometer Distribution



The large dots are the stations involved in the East Indian Ocean Alarm region. PTWC has 14 such alarm regions. PTWC alarms will be triggered by any earthquake with  $M > 5.5$ , and often triggered by much smaller earthquakes.

# Seismic Detection Monitor

STATION COLOR KEY:  $t < 1$  min  $1 < t < 5$  min  $5 < t < 20$  min  $20 < t < 60$  min  $1 < t < 2$  hr





## PTWC General Processes and Procedures for Initial Tsunami Bulletins

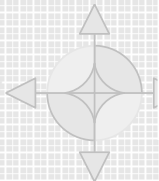
### Refine And Check Results

Interactive  
Hypocenter  
Determination

Interactive  
Magnitude  
Determination

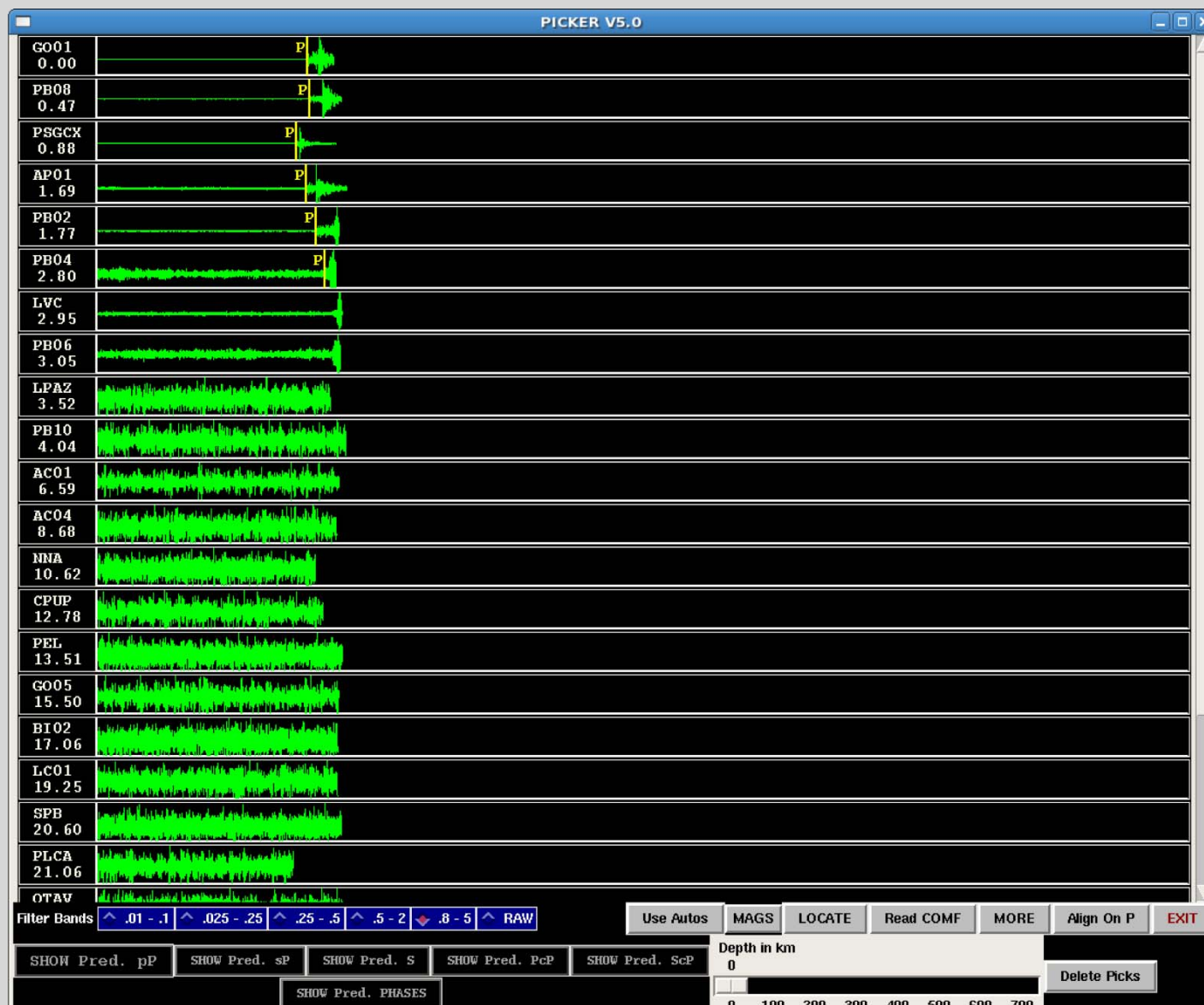
Operator  
Interactive  
Review and  
Refinement

**STEP 5**

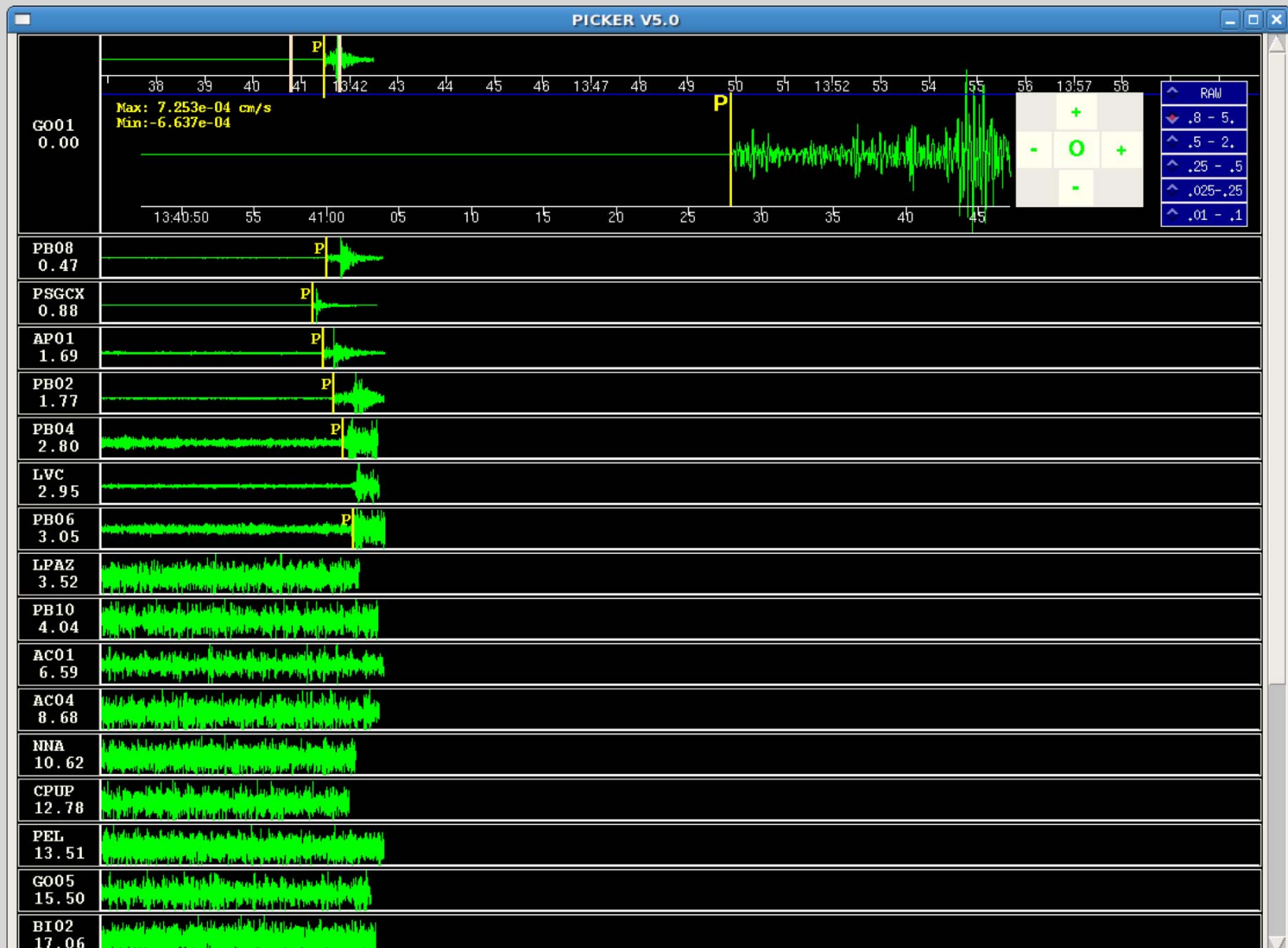




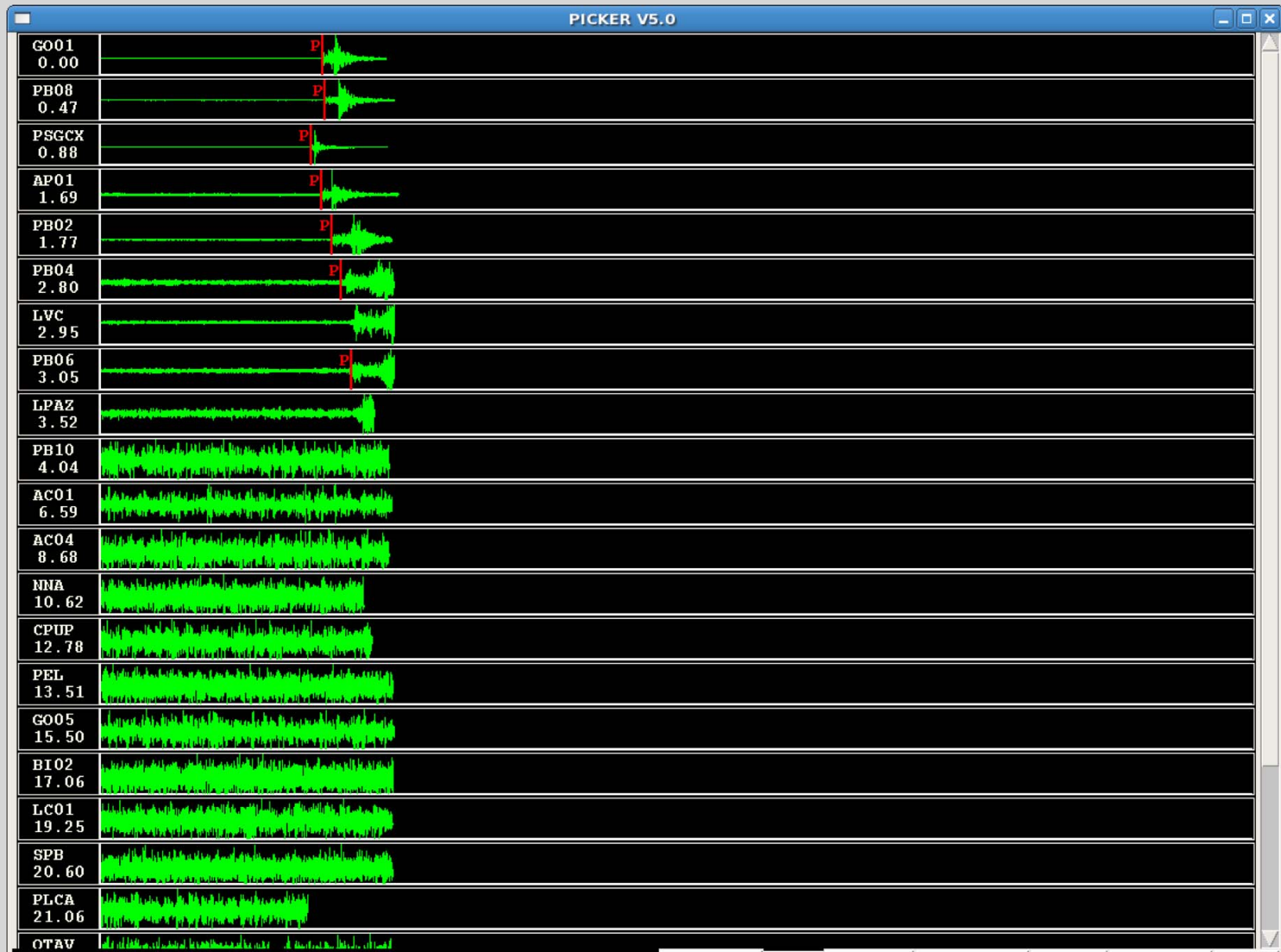
A small earthquake in Chile last night:  
yellow lines are auto picks



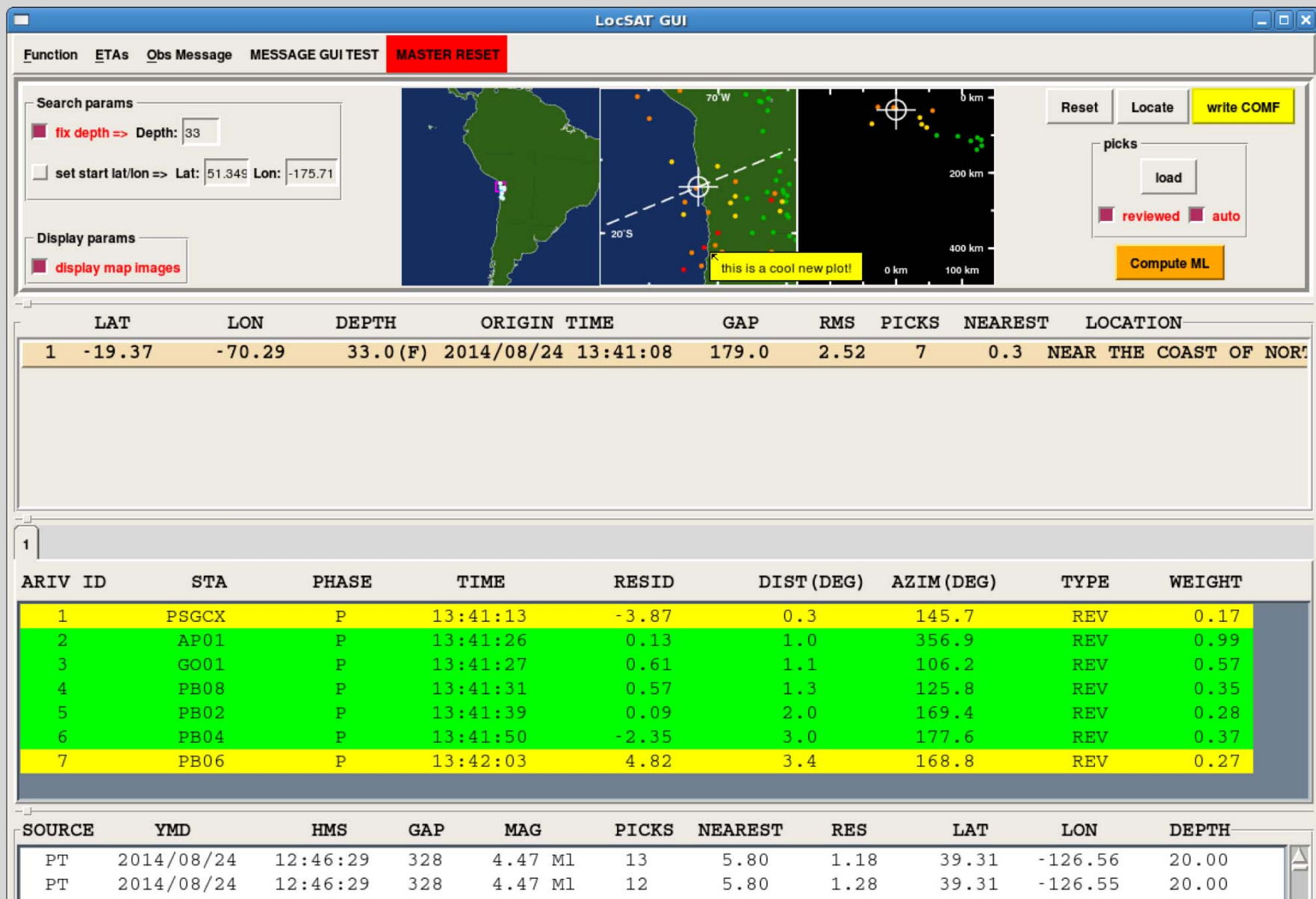
Here I'm reviewing a pick



## Reviewed picks (red lines)

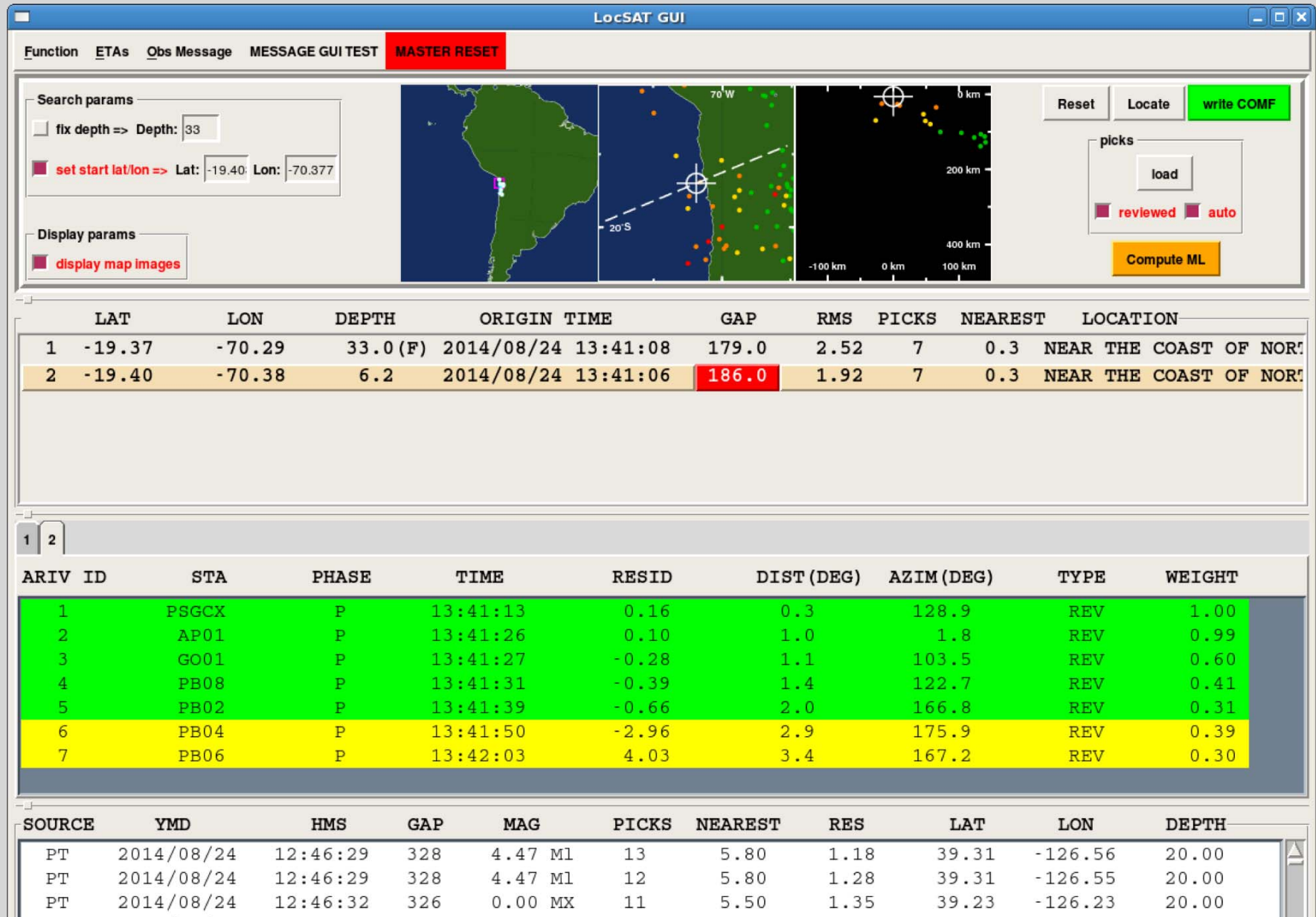


Picks are passed to the location program (Locsat)  
Initially I try a fixed depth (33 km)





Then I allow depth to float to get an improved solution



# Magnitude

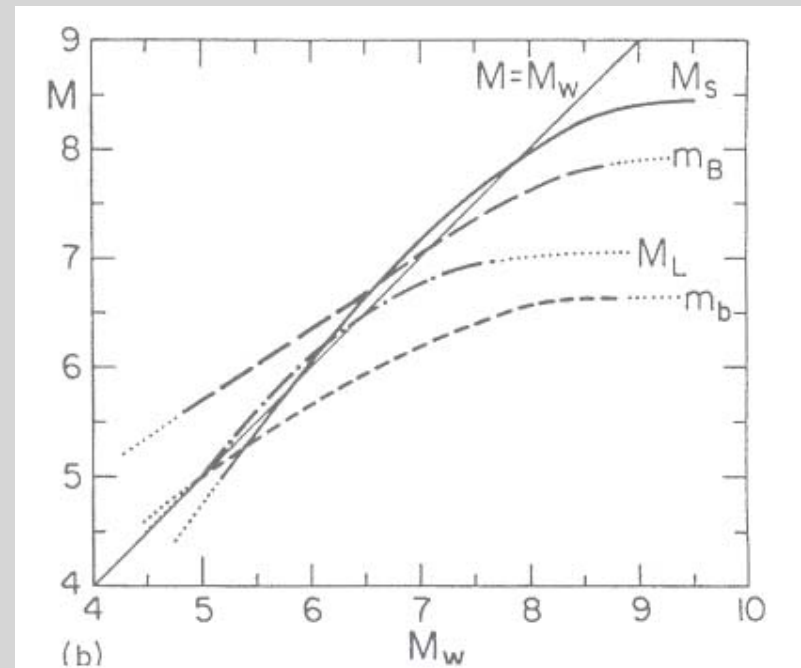
- Once you have the earthquake location you can figure out the magnitude.
- Traditional magnitude measures (Richter magnitude, surface-wave magnitude) are made at too high a frequency to sample the energy from great earthquakes, so those scales saturate. We use broadband magnitude measures (all lumped under the label of “moment magnitude” or  $M_w$ ) for which the problem of saturation is much reduced.
- I’ll talk about two techniques:  $M_{wp}$  (moment magnitude from  $P$ -wave), and the  $W$ -phase determination of moment magnitude.

# Understanding moment magnitude ...or what bumped the Richter scale?



# Saturation

- Traditional magnitude measures (mb, MI, Ms) suffer from saturation.
- Occurs for 2 reasons:
  - The magnitude is calculated for a time window that is less than the duration of the rupture (particularly effects mb)
  - The wavelength of the wave is too short to “see” all of the rupture (effects mb, MI, and Ms)



Kanamori 1983

## $M_{wp}$ (first proposed by Tsuboi, *et al.*, 1995)

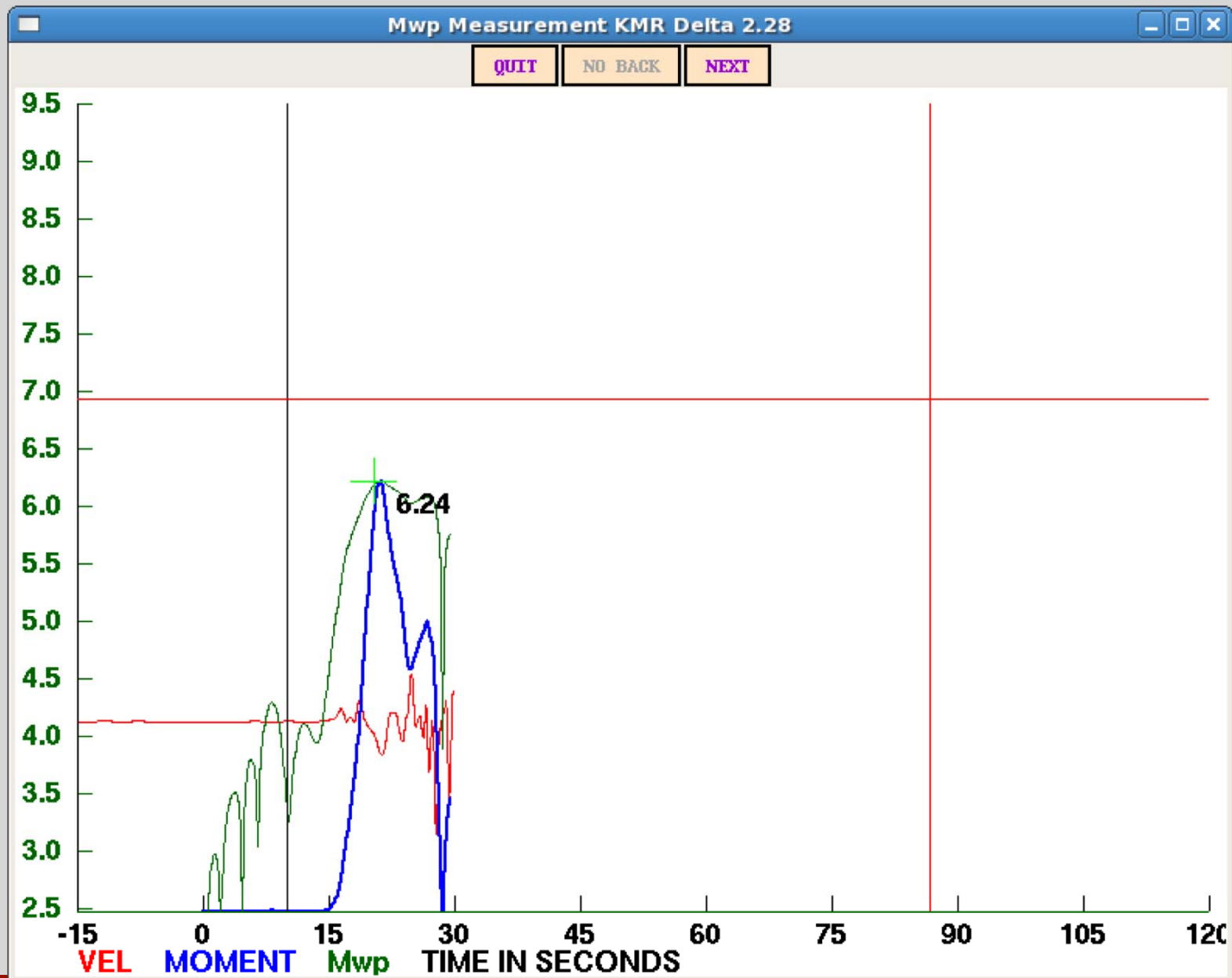
- Problem: Traditional magnitude measurements, while reasonably fast, severely underestimate the size of great earthquakes. Centroid Moment Tensor solutions get the magnitude right, but can take hours—too slow for tsunami warning.
- Suggestion: The displacement record of the first-arriving P waves is a good approximation to the source-time function of the earthquake (i.e., the slip on the fault as a function of time). So integrate the *P*-wave portion of a broadband seismogram from velocity to displacement, integrate again to get seismic moment  $M_0$ . Moment magnitude,  $M_w$ , is then simply obtained from the standard expression

$$M_w = 2/3 \log(M_0) - 10.7$$

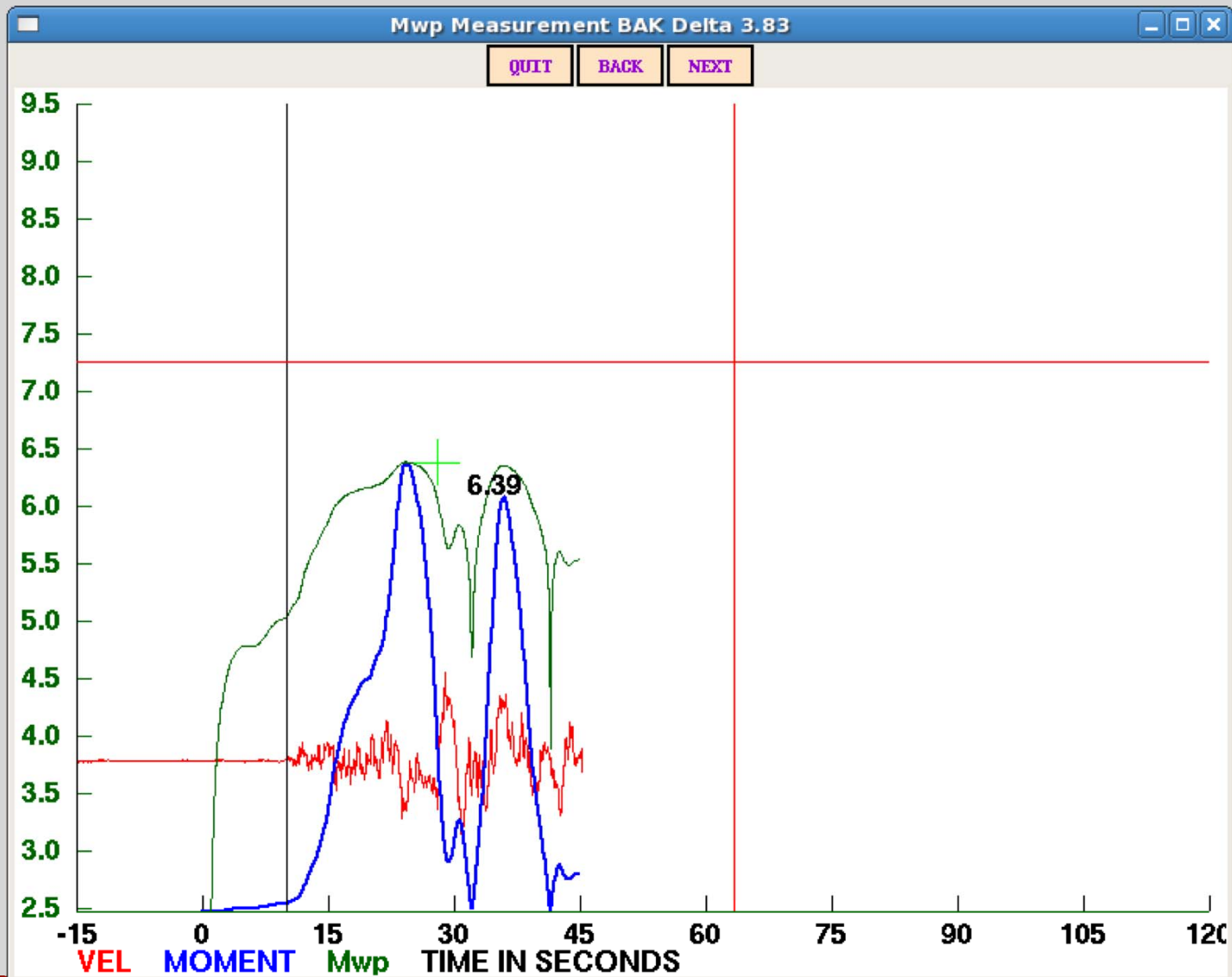
- In practice, the operation is performed on the first 120 seconds of the *P*-wave, or as much as is available before S-waves arrive.



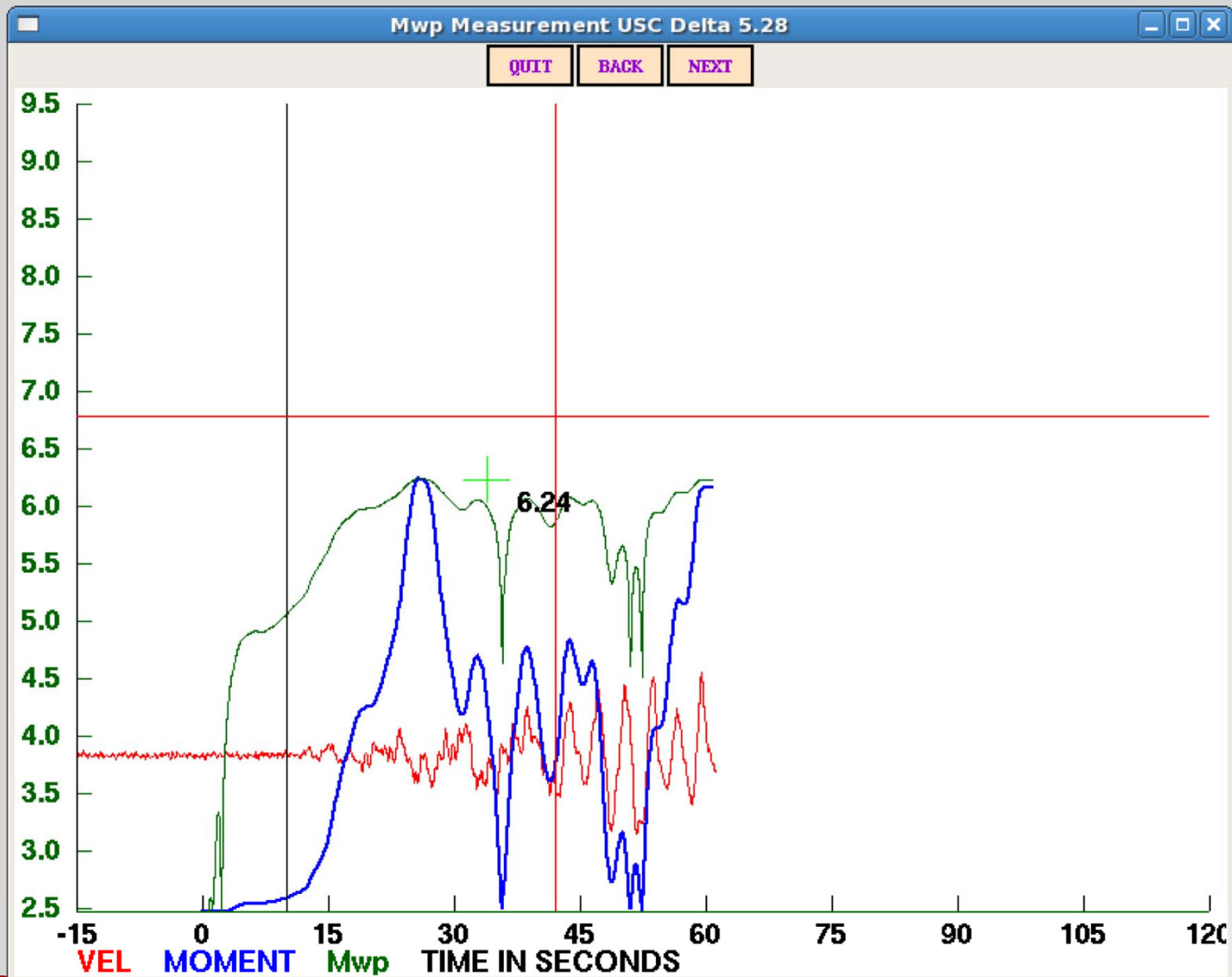
## Example: Yesterday's Napa Earthquake



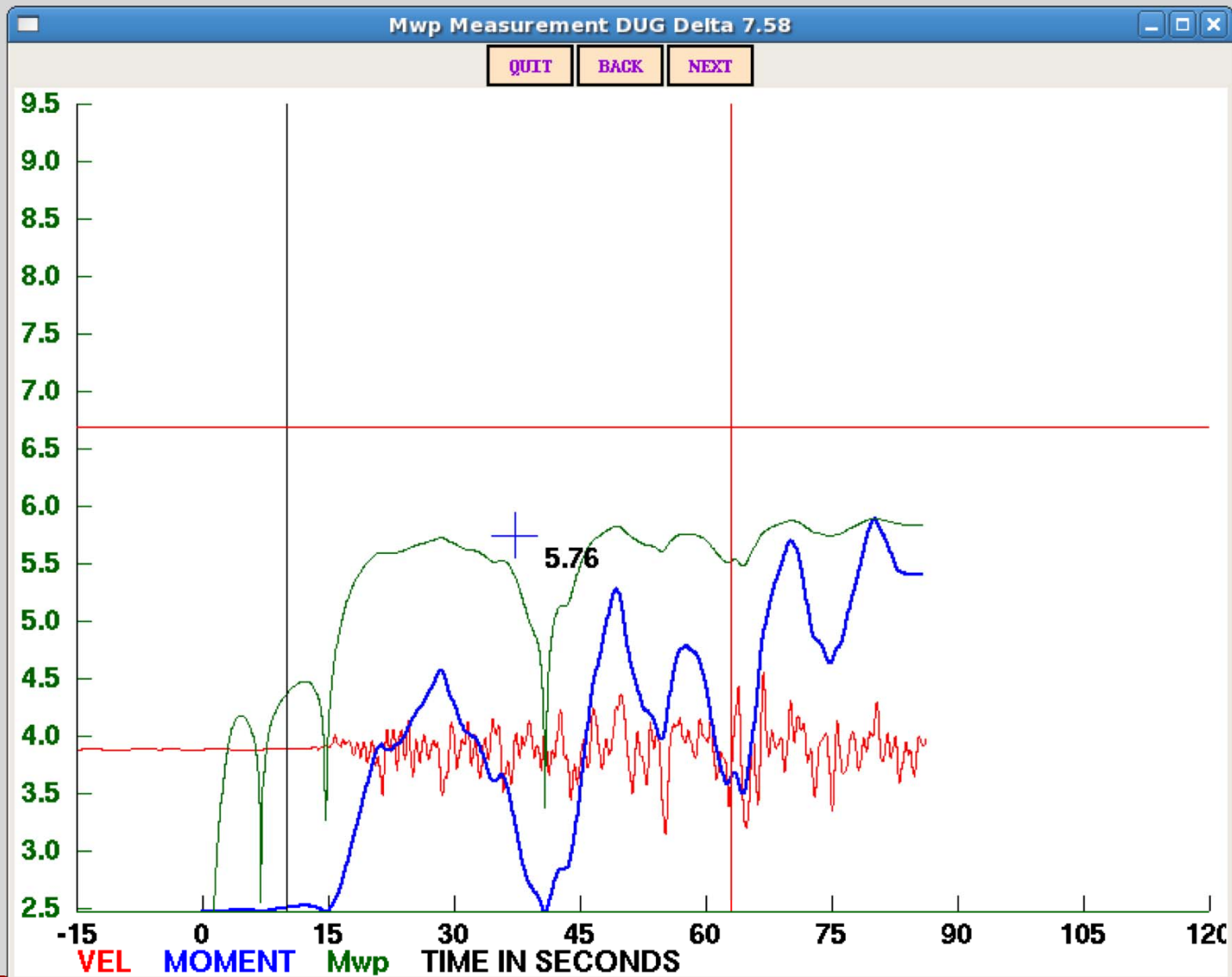
## Example: Yesterday's Napa Earthquake



## Example: Yesterday's Napa Earthquake

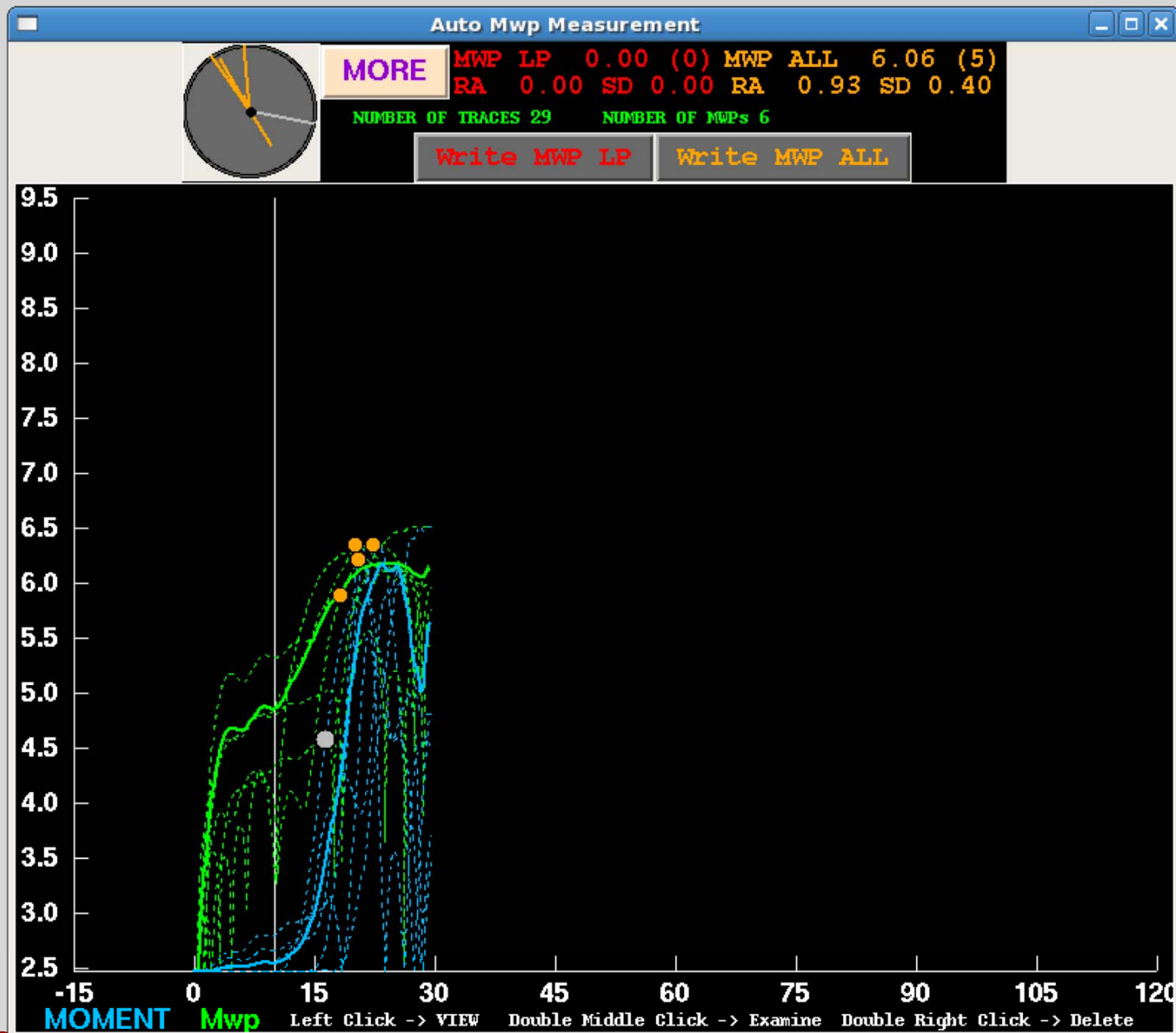


## Example: Yesterday's Napa Earthquake

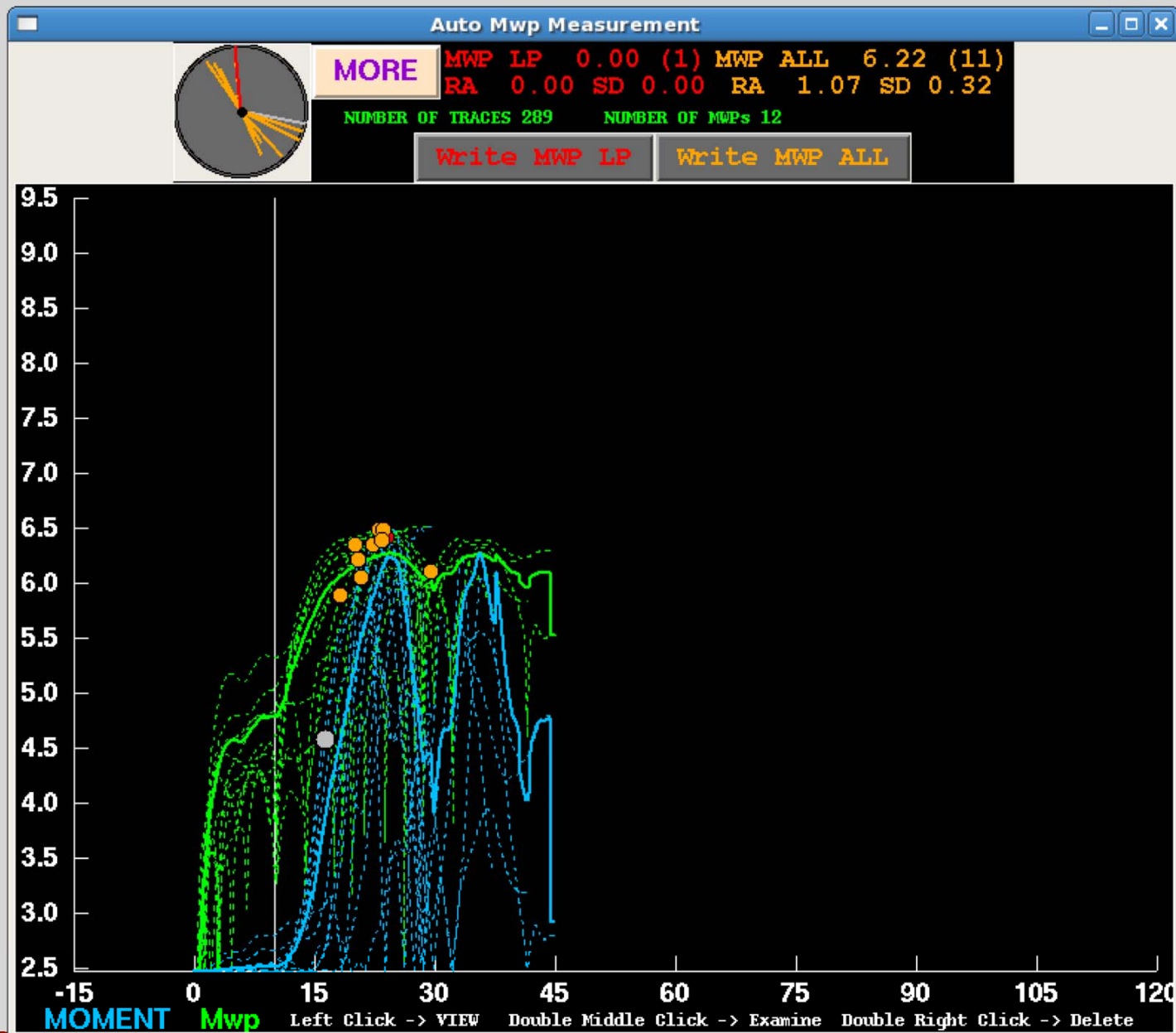




We usually let the  $M_{wp}$  algorithm run automatically



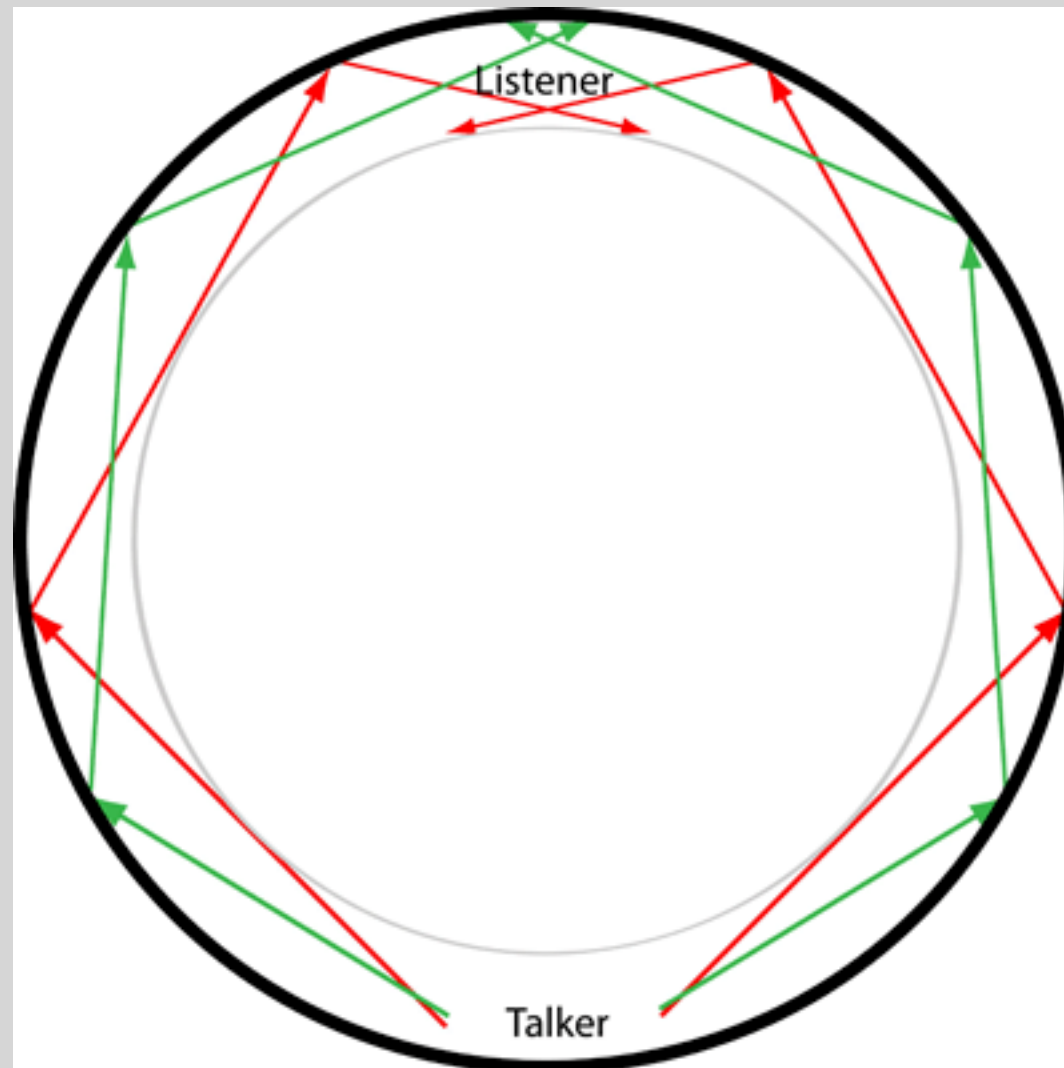
We usually let the  $M_{wp}$  algorithm run automatically



## A diversion: whispering galleries

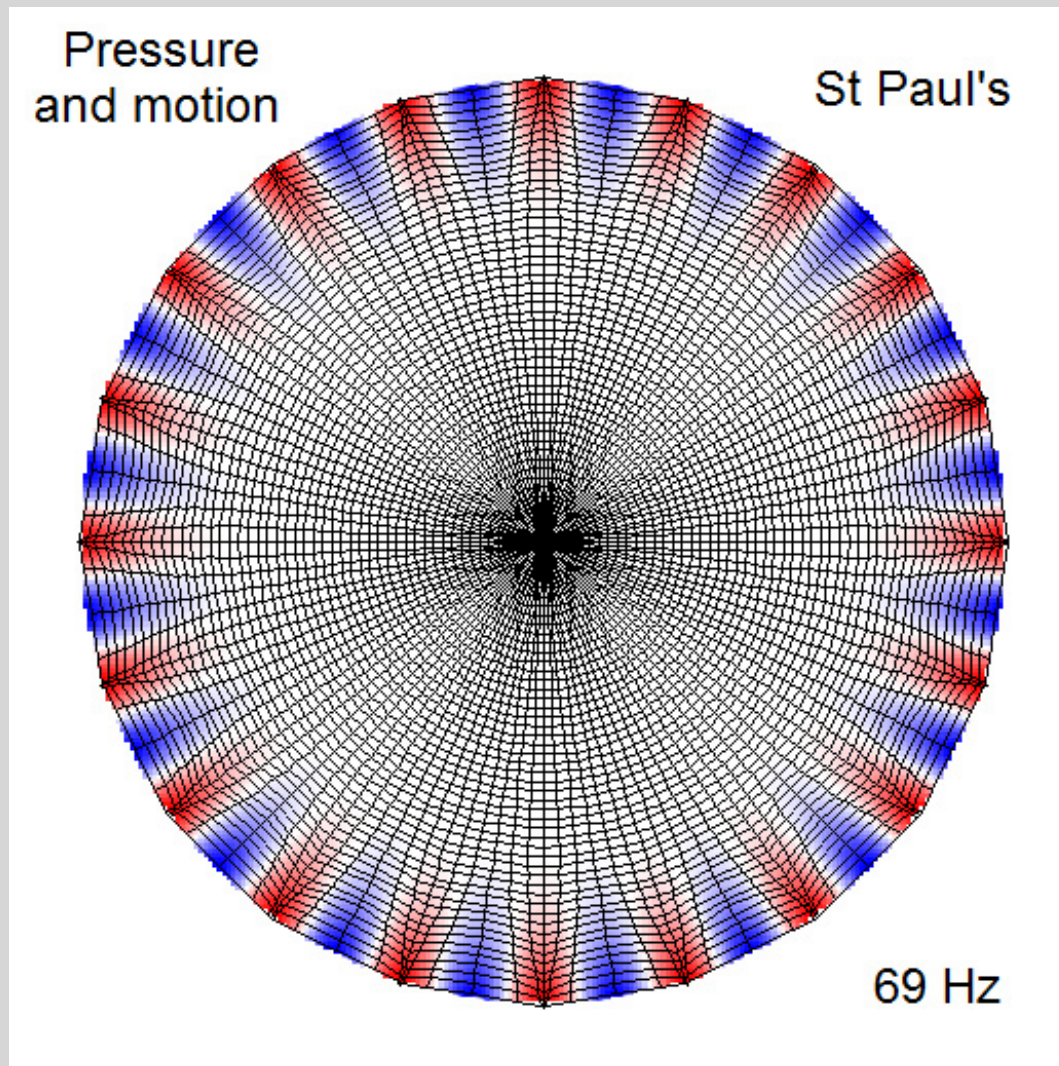


## A diversion: whispering galleries



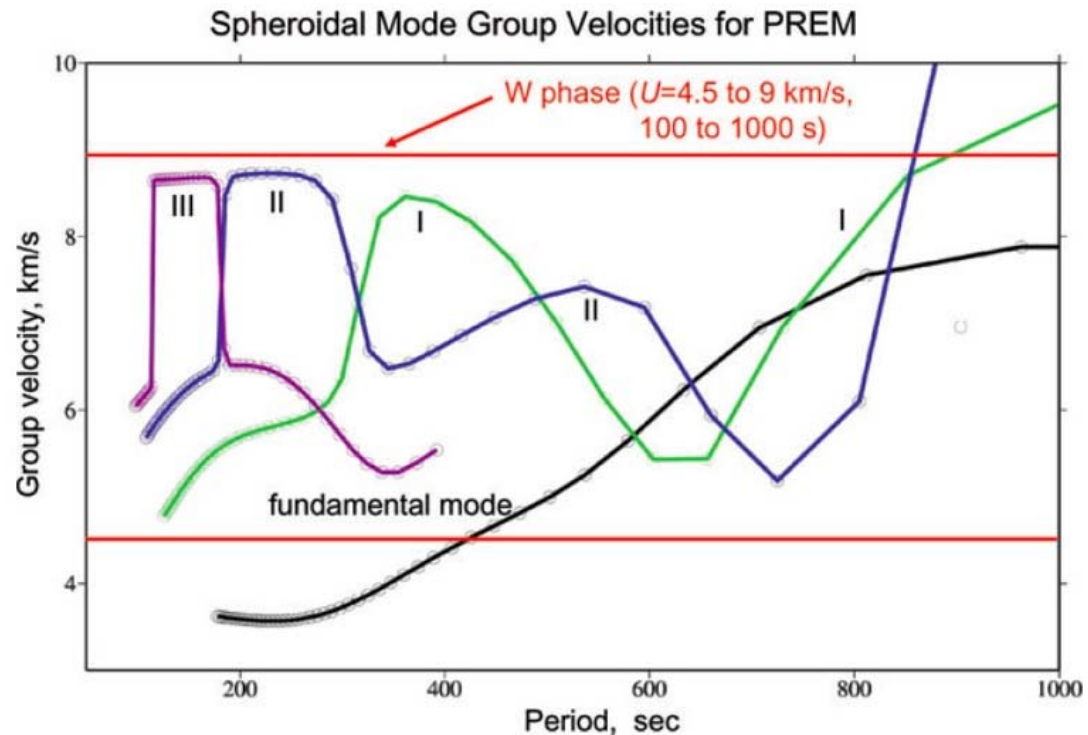
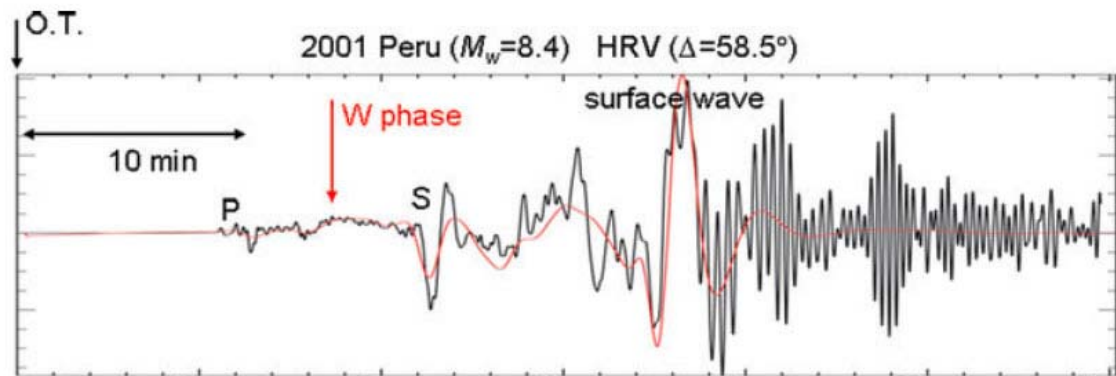


## A diversion: whispering galleries

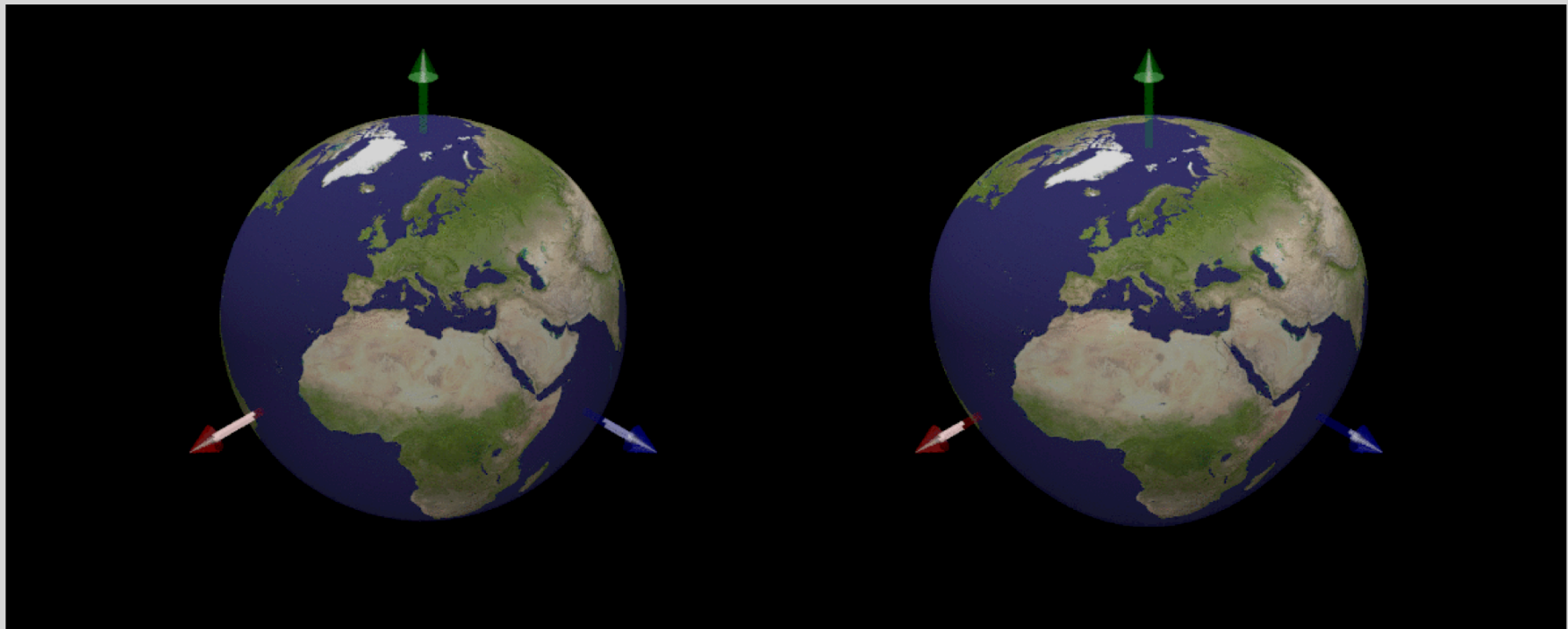


## W-phase

- Ramplike long-period phase between *P* and *S* on displacement seismograms.
- Combination of *P*, *PP*, *PS*, *SP*, *PPP*, etc: whispering gallery phases.
- Combination of spheroidal modes.
- *W*-phase assessment of magnitude currently takes 25 min.



# Spheroidal Normal Modes



$0S_0$

$0S_3$

# W-phase

- Robust magnitude determination: the moment magnitude  $M_w$  from our automatic W-phase code is always correct to within 0.2.
- The algorithm works by solving for the Centroid Moment Tensor of the earthquake, so as a bonus, we get the earthquake mechanism.



# Earthquake Focal Mechanisms

*These describe the direction of slip in an earthquake  
& the orientation of the fault on which it occurs.*





## W-phase solutions from the last few days

Best Double Couple: M0=6.63E+25  
dyn.cm Mw = 6.48  
NP1: Strike= 7 ; Dip=20.3579 ;  
Slip=101  
NP2: Strike=176 ; Dip=70.0154 ;  
Slip= 86

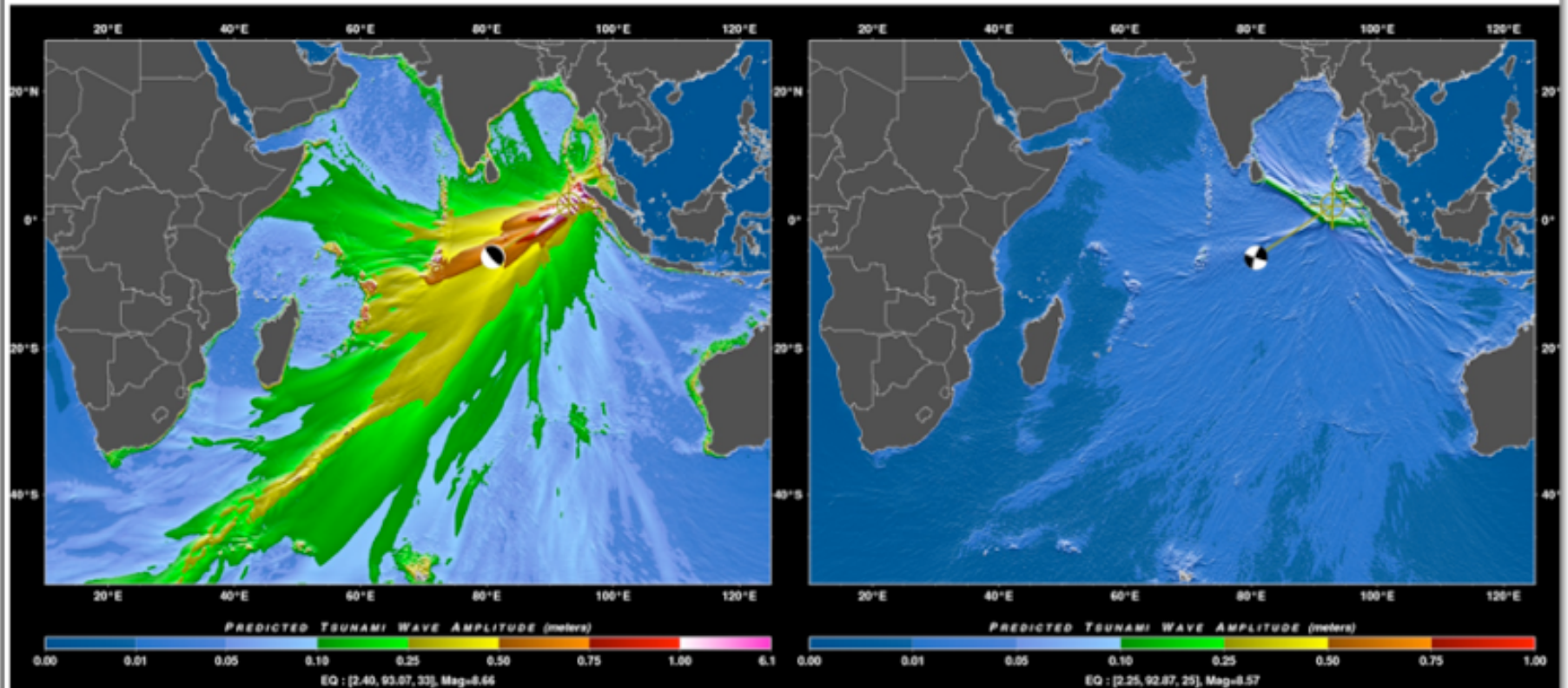
[illegible]

Best Double Couple: M0=2.58E+26  
dyn.cm Mw = 6.87  
NP1: Strike=195 ; Dip=44.4631 ;  
Slip=-68  
NP2: Strike=346 ; Dip=49.4481 ;  
Slip=-110

[illegible]

# The mechanism matters

## PTWC Real Time Forecast of Tsunamis (RIFT)

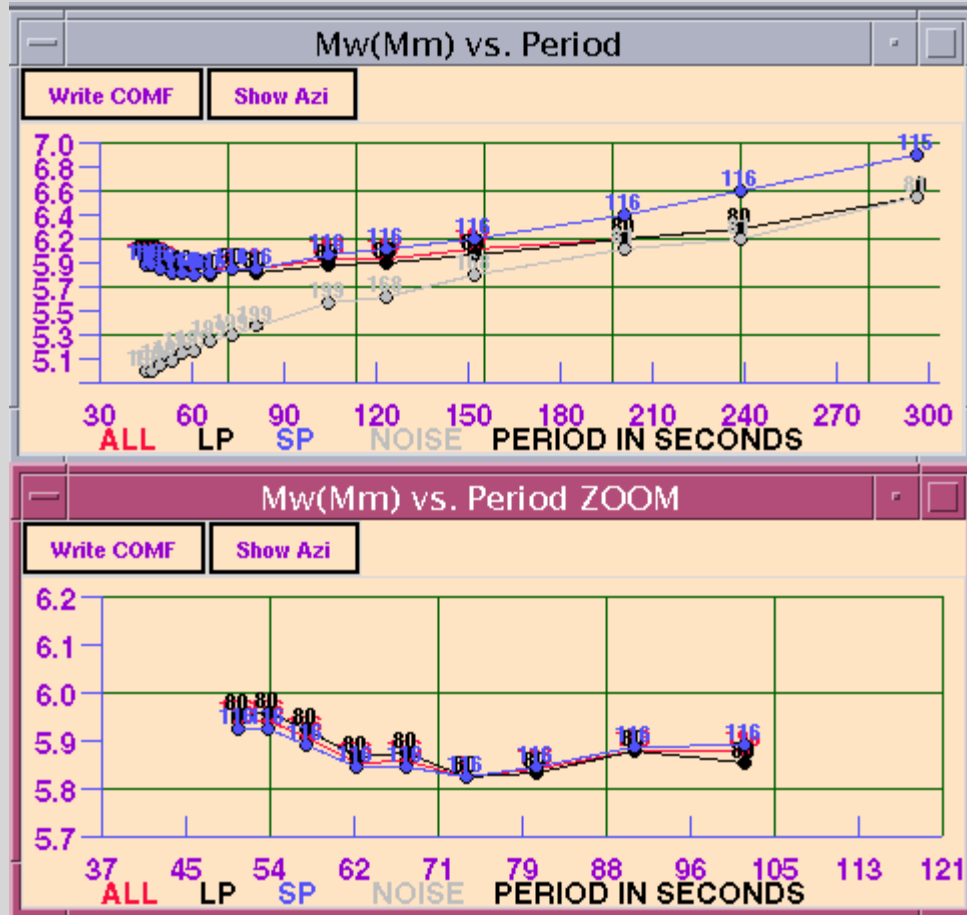


Based on shallow thrust assumption  
Mwp=8.66 (initial PTWC Magnitude)

USGS (Gavin Hayes )  
Wphase CMT: Mw=8.57 (Strike-slip)

The initial PTWC magnitude is larger than that of the USGS CMT magnitude, but the major difference between the two runs are focal mechanisms: thrust vs. strike-slip.

# Mantle magnitude $M_m$ Largely supplanted by $W$ -phase



Mantle Magnitude

EXIT 13:11:59 01/10

|      |      |       |      |      |
|------|------|-------|------|------|
| CTAO | DONE | ----  | 6.37 | 6.85 |
| RCBR | DONE | ----  | 4.76 | 5.78 |
| GUMO | DONE | ----  | 6.09 | 6.66 |
| WAKE | DONE | ----  | 6.10 | 6.67 |
| KAPI | TIME | NA    | ---- | ---- |
| MBWA | GAP  | NA    | ---- | ---- |
| TAU  | DONE | ----  | 6.31 | 6.80 |
| SNZO | DONE | ----  | 4.39 | 5.53 |
| TATO | DONE | ----  | 6.04 | 6.62 |
| MIDW | DONE | ----  | 6.42 | 6.88 |
| NWAO | DONE | ----  | 5.88 | 6.52 |
| MAJO | DONE | ----  | 5.90 | 6.53 |
| INCN | PEND | 00:00 | ---- | ---- |
| YSS  | PEND | 00:00 | ---- | ---- |
| KIP  | DONE | ----  | 6.32 | 6.81 |
| POHA | DONE | ----  | 6.47 | 6.91 |
| COCO | PEND | 00:00 | ---- | ---- |
| BJT  | PEND | 00:00 | ---- | ---- |
| PET  | DONE | ----  | 6.19 | 6.73 |
| SMY  | PEND | 00:00 | ---- | ---- |
| HIA  | PEND | 00:00 | ---- | ---- |
| ADK  | DONE | ----  | 6.31 | 6.81 |
| MA2  | DONE | ----  | 6.05 | 6.63 |
| ULN  | DONE | ----  | 6.04 | 6.63 |
| UNV  | DONE | ----  | 6.99 | 7.26 |
| YAK  | PEND | 00:00 | ---- | ---- |
| TLY  | DONE | ----  | 6.24 | 6.76 |
| CCM  | DONE | ----  | 6.07 | 6.64 |
| VNDA | PEND | 00:00 | ---- | ---- |
| SBA  | DONE | ----  | 6.10 | 6.67 |

SLEEPING 08 Mean Mm is: 6.12(24) Mw is: 6.68 PRINT

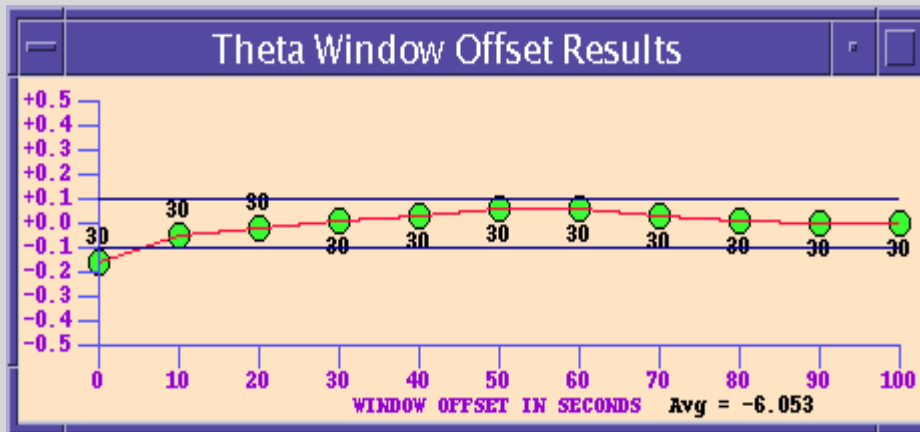


# SLOW EARTHQUAKES

$$\text{Theta} = \log_{10} ( E_r / M_o )$$

$E_r$  is energy carried by high frequency P-waves and  $M_o$  is seismic Moment.

Ratio of energy contained by high frequency P-waves to energy released by EQ.



Slow, or tsunami earthquake!

| Energy Discriminant          |        |       |       |
|------------------------------|--------|-------|-------|
| EXIT 01:05:19 09/02 Mw: 7.20 |        |       |       |
| STA                          | STATUS | ETA   | THETA |
| HIA                          | PEND   | 00:00 | ----  |
| KDAK                         | DONE   | ----- | -6.23 |
| YAK                          | DONE   | ----- | -5.32 |
| ULN                          | DONE   | ----- | -5.37 |
| TNA                          | DONE   | ----- | -6.37 |
| BILL                         | PEND   | 00:00 | ----  |
| RPN                          | PEND   | 00:00 | ----  |
| PMR                          | DONE   | ----- | -5.81 |
| EYAK                         | DONE   | ----- | -5.86 |
| DIV                          | PEND   | 00:00 | ----  |
| MCK                          | DONE   | ----- | -5.86 |
| LSA                          | PEND   | 00:00 | ----  |
| JCC                          | DONE   | ----- | -5.94 |
| SAO                          | DONE   | ----- | -6.06 |
| SIT                          | DONE   | ----- | -6.14 |
| COLA                         | DONE   | ----- | -5.74 |
| SNCC                         | PEND   | 00:00 | ----  |
| PKD                          | DONE   | ----- | -6.03 |
| CRAG                         | DONE   | ----- | -6.06 |
| WDC                          | DONE   | ----- | -6.00 |
| YBH                          | DONE   | ----- | -6.10 |
| ORV                          | DONE   | ----- | -6.30 |
| CMB                          | DONE   | ----- | -6.18 |
| OSI                          | PEND   | 00:00 | ----  |
| COR                          | DONE   | ----- | -5.94 |
| PAF                          | GAP    | NA    | ----  |
| SKAG                         | DONE   | ----- | -6.09 |
| PAS                          | PEND   | 00:00 | ----  |
| KCC                          | DONE   | ----- | REM   |
| ISA                          | PEND   | 00:00 | ----  |

FINISHED

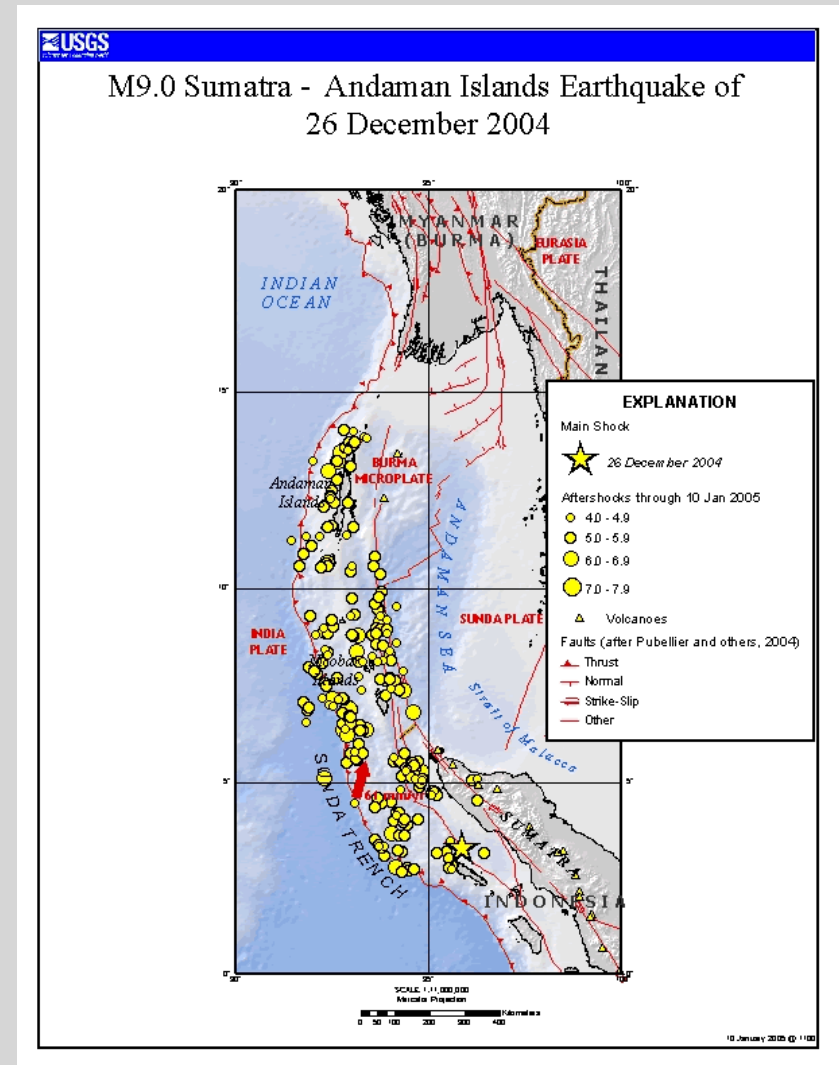
Mean Theta is: -5.83(43)

PRINT



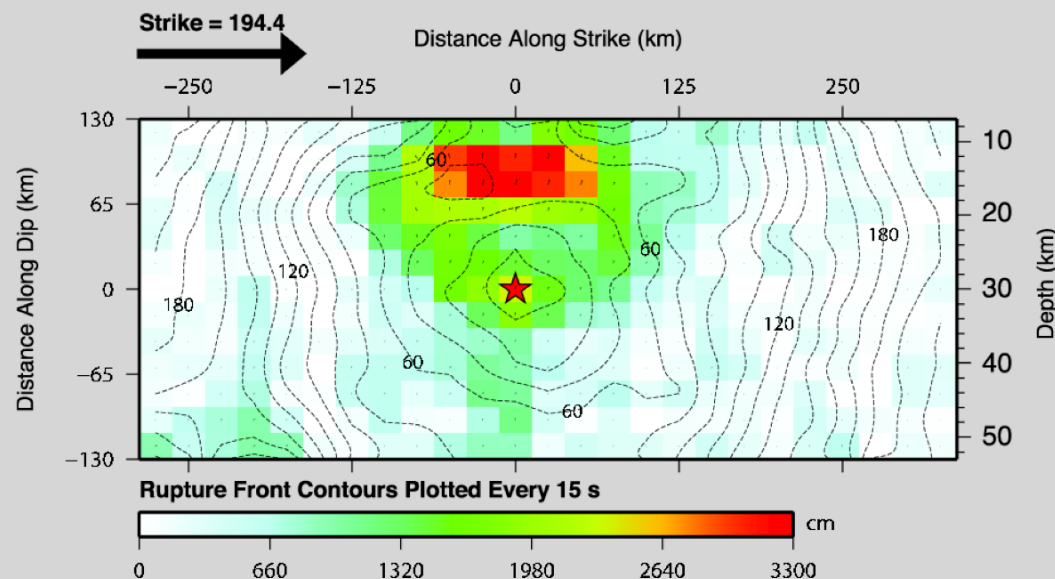
# Weaknesses we know about

- Where did the earthquake go? We can deduce epicenter and magnitude in a few minutes, but what happened to the rupture? Fortunately, all post-2004 great earthquakes have had bilateral rupture.
- Fault extent. We need a rapid finite fault procedure (i.e, a description of how rupture is distributed over the fault plane).
- We are not really equipped to handle non-earthquake sources.





# Our seismological future



- Fast (crude) finite fault analysis.
- Faster *W*-phase and other automatic techniques (Rivera, Kanamori, Duputel, Lomax, etc.)
- High-rate GPS for rapid source determination.
- We should be able to get  $M_w$  and mechanism within 10 minutes.

Mahalo

